

# Cloud Model Upgrade to the GIFTS Fast Radiative Transfer Model

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## 1 INTRODUCTION

The GIFTS fast radiative transfer model (GIFTSFRTE) computes simulated top-of-atmosphere radiances and/or brightness temperatures for the GIFTS instrument. On 7 January 2003, a fast radiative transfer code for cirrus ice clouds and liquid phase clouds [9] was delivered to SSEC. The purpose of this document is to detail the installation, testing and verification of this code and the changes made to the GIFTS fast radiative transfer model to facilitate the incorporation of the new cloud model. Throughout this document we shall refer to the cloud model in GIFTSFRTE prior to this release as the Old cloud model. The current, newly implemented model will be referred to as Yang:2003.

## 2 DELIVERED CODE

The new cloud model (Yang:2003) was created under contract to SSEC by Ping Yang and Heli Wei at Texas A&M University. Yang:2003 was integrated by the Heli Wei into GIFTSFRTE and supercedes the Old cloud model in the code development branch `/cvs/devel/Model/giftsrte/src/giftsfrte.f`. Upon delivery some code changes were made at SSEC to allow successful compilation with GNU `g77` under the LINUX operating system. Specifically, some F90 MODULE definitions were replaced with an include file (`texcloud.i`) comprising certain parameters and a common block (`/wei_common/`). These changes were made expediently to allow the code to be tested in the presence of one of the code developers, namely Heli Wei, during his short visit to SSEC.

Fig. 1 shows the common blocks named in the source modules that make up the GIFTS fast radiative transfer model. Fig. 2 shows the call tree for the revised GIFTSFRTE, as it was delivered on 7 January 2003, but after the minor code modifications indicated in the preceding paragraph. In Fig. 2, function `radwnplank` is equivalent to `wnplan` (in source file `wnplan.f`) except that `radwnplank` accepts a single precision wavenumber whereas `wnplan` accepts a double precision wavenumber. Text files `parameters.txt`, `cld.par` and `testout.dat` are hard-coded into the source file `giftsfrte.f`, as are binary files `wvnLW.bin` and `wvnSMW.bin`. Binary files `test.bin`, `outLW.bin` and `outSMW.bin` are listed in the parameter file `parameters.txt`, and the text file `usstandout.dat` is named in the cloud parameter file `cld.par`. The source module `gifts_tran_od` is `gifts_tran` modified to return optical depths.

---

<u>common block</u>	<u>source file</u>	<u>routines</u>
<code>/wei_common/</code>	<code>texcloud.i</code>	<code>main</code> <code>read_fitted_database</code> <code>interpolation_rt</code>
<code>/gifts_vn/</code>	<code>giftsfrte.f</code> <code>vn_gifts.f</code>	<code>main</code> <code>vn_gifts</code>
<code>/taudwo/</code>	<code>giftsfrte.f</code> <code>gifts_tran_od.f</code>	<code>main</code> <code>gifts_tran_od</code>
<code>/atmstd/</code>	<code>irt_sub_101.f</code> <code>gifts_tran_od.f</code>	<code>reference_atmosphere (block data)</code> <code>gifts_tran_od</code>
<code>/gifts_be/</code>	<code>vn_gifts.f</code>	<code>vn_gifts</code>

---

Figure 1: COMMON blocks in GIFTS forward model with Yang:2003 cloud model (as delivered).

---

source file	call tree	inputs from	outputs to
giftsfrte.f	main   (texcloud.i)	STDIN parameters.txt (test.bin)  cld.par (usstandout.dat)	STDOUT testout.dat (outLW.bin) (outSMW.bin) wvnLW.bin wvnSMW.bin
subroutines.f	+--read_fitted_database   (texcloud.i)	fitedcoeRc.dat fitedcoeRw.dat fitedcoeTc.dat fitedcoeTw.dat	
	+--topbt		
	+--interpolation_rt     (texcloud.i)		
	+--sumfit		
	+--radwnplank		
	+--radtolayer		
	+--radwnplank		
wnbrit.f	+--wnbrit		
	+--wnbrit		
gifts_tran_od.f	+--gifts_tran_od	giftsfr1.dry giftsfr1.ozo giftsfr1.wco giftsfr1.wtl giftsfr1.wts giftsfr2.dry giftsfr2.ozo giftsfr2.wco giftsfr2.wtl giftsfr2.wts	
irt_sub_101.f	+--conpir     +--gphite		
	+--taudoc		
	+--calpir		
	+--tauwtr		
lsurface.f	+--lsurface		
wmrad.f	+--wmrad		
vn_gifts.f	+--vn_gifts		

---

Figure 2: Call tree for GIFTS forward model with Yang:2003 cloud model (as delivered). The include file `texcloud.i` is included in the program units indicated. Files in brackets, e.g.(`test.bin`), are named in input files.



### 3 GIFTS FAST MODEL UPGRADE

After delivery, changes were made to prepare GIFTSEFRTE for production release. `giftsfrte.f` was split into two source code modules, named `giftsfrte_main.f` and `giftsfrte.f`. The main program absorbs the task formerly handled by subroutine `read_fitted_database`, however the coefficient files to be read are left as they were delivered, i.e. as text files, but have been concatenated into a single file, namely `clouds.dat`. The remaining tasks are performed by subroutine GIFTSEFRTE. This separation of the fast model into a main program (performing input and output), and a subprogram (responsible, with its subsidiary routines, for the science outcome) is a design goal for future GIFTS code development. This source code structure simplifies the preparation of pipeline science-only modules for cluster processing, a future requirement for GIFTS real-time data processing (not covered in this document).

The new call tree is shown in Fig. 3. Subprogram units `topbt`, `interpolation_rt`, `sumfit` and `rادتolayer` (and the now redundant subroutine `read_fitted_database`) have been combined into a single file, that being `texas.cloud_model.f`.

Legacy codes `wnbrit`, `wnmrad`, `lsurface` and the utility routines in `irt_sub_101.f` have not been modified (other than to embed revision information in their compiled object modules). The GIFTS fast transmittance code, `gifts_tran_od`, has been replaced with `gifts_tran` from the main code development branch.<sup>1</sup> `gifts_tran` has also been modified to access the regression coefficients it requires in directories other than the current working directory.

Inputs are from coefficient files, atmospheric profile files from the mesoscale model MM5 [3] and from standard input, typically in the form of a redirected parameter file. The parameter file contains the paths to the regression coefficient files (the wildcard expansion of `{giftsfr?.*}`) and to the cloud parameterization (`clouds.dat`) together with a range of processing options. An example GIFTSEFRTE parameter file can be found in Appendix C.

---

<sup>1</sup>`gifts_tran_od` is an offshoot developed to satisfy a “one-off” request that was never checked back into the code repository.



Figure 3: Call tree for upgraded GIFTS fast radiative transfer model (GIFTSFRTE). Inputs are from coefficient files, MM5 atmospheric profile files and from standard input, typically in the form of a redirected parameter file.

The include file previously named `texcloud.i` has been renamed to `giftsfrte.i` and the common block nominated therein is now called `/cld_common/`. In spite of a preference in modern FORTRAN code development to avoid the use of COMMON blocks, they have not been eliminated in this code release in the interests of delivering the modified code in a reasonable time and without introducing errors into existing, stable code (e.g. `gifts_tran()` which references common blocks `/taudwo/` and `/atmstd/`). The COMMON blocks in this code release are listed in Fig. 4. A GNU makefile (Fig. 5) generates the executable called `giftsfrte` (not `giftsfrte.main`).

---

common block	source file	routines
-----	-----	-----
<code>/cld_common/</code>	<code>giftsfrte.i</code>	<code>main</code> <code>interpolation_rt</code>
<code>/gifts_vn/</code>	<code>giftsfrte.f</code> <code>vn_gifts.f</code>	<code>giftsfrte</code> <code>vn_gifts</code>
<code>/taudwo/</code>	<code>giftsfrte.f</code> <code>gifts_tran.f</code>	<code>giftsfrte</code> <code>gifts_tran</code>
<code>/atmstd/</code>	<code>irt_sub_101.f</code> <code>gifts_tran.f</code>	<code>reference_atmosphere (block data)</code> <code>gifts_tran</code>
<code>/gifts_be/</code>	<code>vn_gifts.f</code>	<code>vn_gifts</code>

---

Figure 4: COMMON blocks in the upgraded GIFTS fast radiative transfer model GIFTSFRTE.

---

```
#
# $Id: Makefile,v 1.4 2003/06/06 20:31:31 rayg Exp $
#
# Description:
# Makefile for the GIFTS fast radiative transfer model with clouds
#
FC      = g77
FFLAGS = -Wall -O3 ${CFLAGS}
LFLAGS = -static
LIBS    =
BASE    = giftsfrte
EXEC    = ../bin/giftsfrte
INCL    = ${BASE}.i
SRCS    = ${BASE}_main.f ${BASE}.f gifts_tran.f irtsb_101.f vn_gifts.f \
          lsurface.f wnbrit.f wnmrad.f wnplan.f texas_cloud_model.f atmosrad.f \
          strcompress.f
OBJS    = ${SRCS:.f=.o}

${EXEC}: ${OBJS}
${FC} ${OBJS} ${LFLAGS} ${LIBS} -o ${EXEC}

${OBJS}: ${INCL}

${INCL}:

clean:
/bin/rm -f ${EXEC} ${OBJS}
```

---

Figure 5: Makefile for GIFTSFRTE.

## 4 GIFTS DATA CUBES

GIFTS data cubes comprise a  $128 \times 128$  array, with a third dimension that contains spectral information (such as radiance, brightness temperature, transmittance or optical depth). A  $128 \times 128$  array of atmospheric profiles and surface conditions may also be considered as a GIFTS data cube, these data are the output of MM5 model executions (reformatted for use with GIFTS fast model simulations) and one of the inputs to GIFTSFRTE. Within the  $128 \times 128$  data arrays, row and column pixels are defined to be on the interval  $[-64:64]$  with row 0 and column 0 skipped.

Some example contour plots drawn from MM5 data are shown in the next few figures. Fig. 6 shows the surface heights for GIFTS data cube 06122002.2200.2.3. The data cube is centered at approximately  $34.51^\circ$  N,  $86.82^\circ$  W, each pixel is 4 km square.

Fig. 7 and Fig. 8 show the total column liquid water and ice, respectively, for GIFTS data cube 06122002.2200.2.3. There is an accumulation of column liquid water in

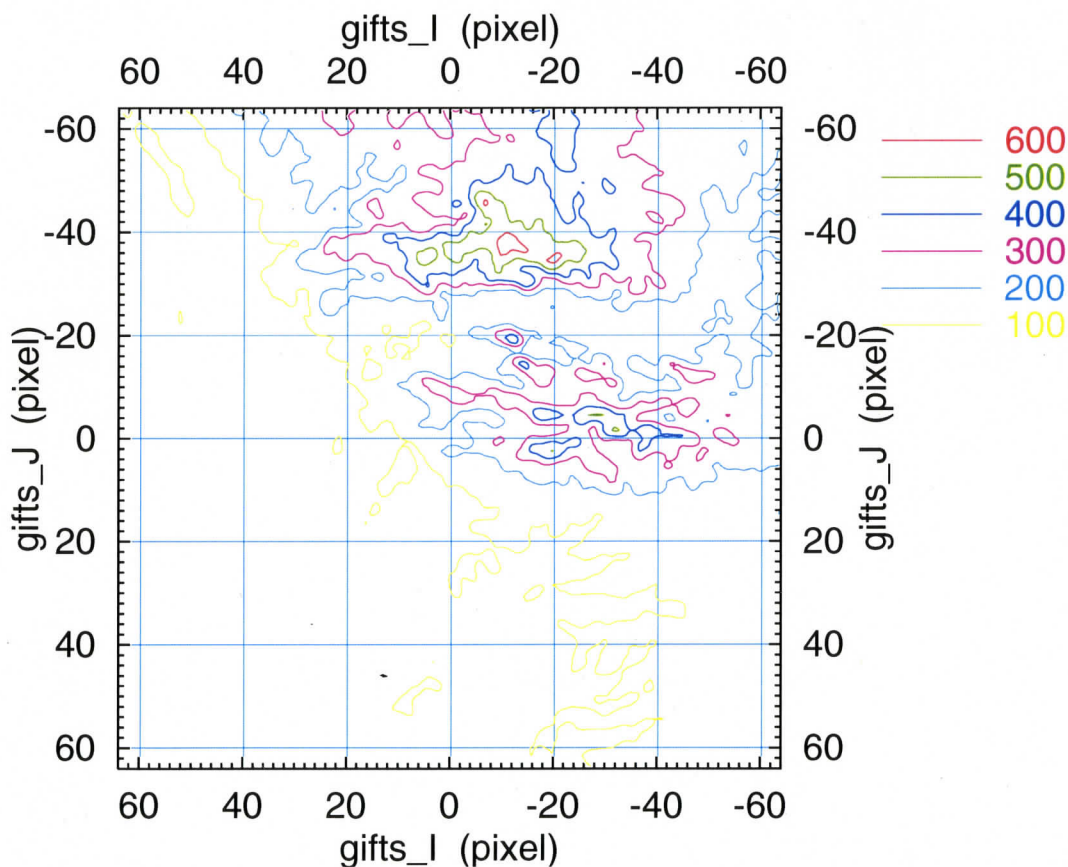


Figure 6: Surface height (m) contour plot for GIFTS data cube 06122002.2200.2.3.

the vicinity of  $(-31, -32)$ .

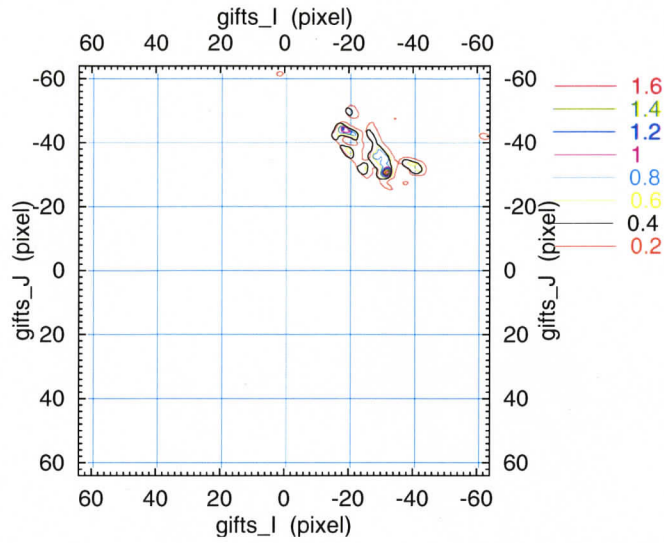


Figure 7: Liquid water path ( $\text{g/cm}^2$ ) contour plot for GIFTS data cube 06122002.2200.2.3.

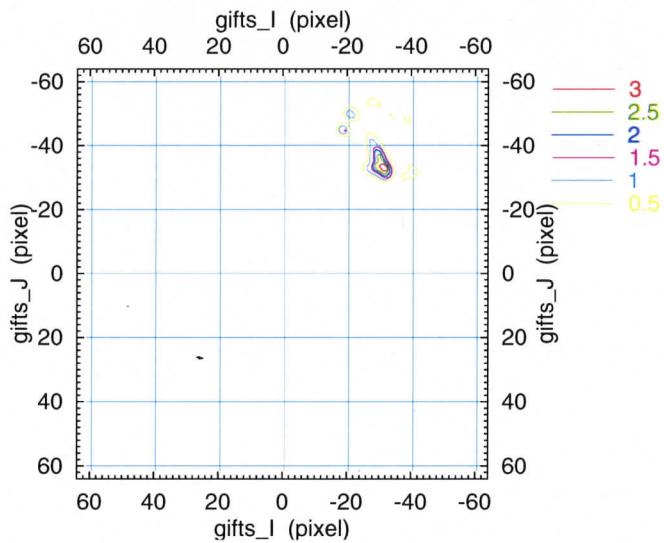


Figure 8: Ice water paths ( $\text{g/cm}^2$ ) contour plot for GIFTS data cube 06122002.2200.2.3.

### 4.1 Atmospheric Profile Data

Atmospheric profile data are represented as 312 single precision floating point numbers in a binary record; 16384 ( $128 \times 128$ ) records constitute a GIFTS atmospheric profile data cube. The first 303 values per record are ordered as 101 temperatures (K), 101 water vapour concentrations (g/kg) and 101 ozone concentrations (ppmv). The ordering is lowest pressure to highest pressure. The 101 values are matched to 101 standard pressure levels to be found as data in source file `giftsfrte_main.f`. The remaining 9 values are, in this order, liquid water path ( $\text{g/m}^2$ ), ice water path ( $\text{g/m}^2$ ), surface skin temperature (K), surface altitude (m), latitude (deg +N), longitude (deg +E), pressure level where liquid condensate is located (hPa), pressure level where ice condensate is located (hPa) and, finally, surface pressure (hPa). Fig. 9 shows an example atmospheric profile together with values of these nine non-profile data quantities. The brightness temperatures that GIFTSFRTE computes for this cloud-free atmosphere are shown in Fig. 10 (LW band) and Fig. 11 (SMW band).

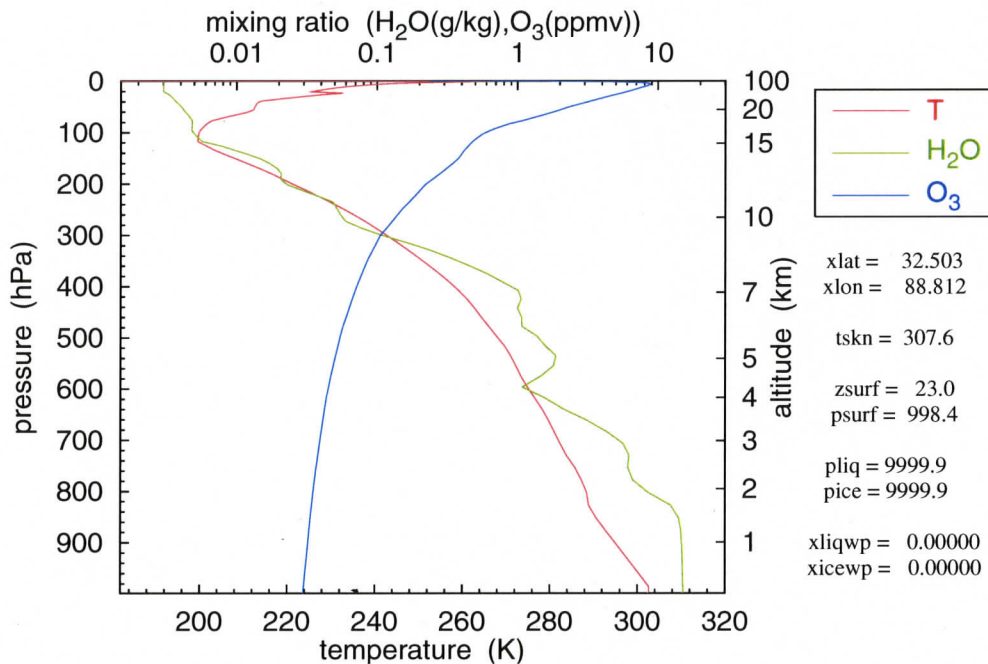


Figure 9: Example atmospheric profile. The nine non-profile quantities are listed to the right of the plotted profile data. In this example the pressure levels of liquid water and ice condensate, `pliq` and `pice` respectively, have been assigned flag values. Liquid water and ice water paths, `xliqwp` and `xicewp` respectively, are zero — i.e. this is a “clear” atmospheric profile.

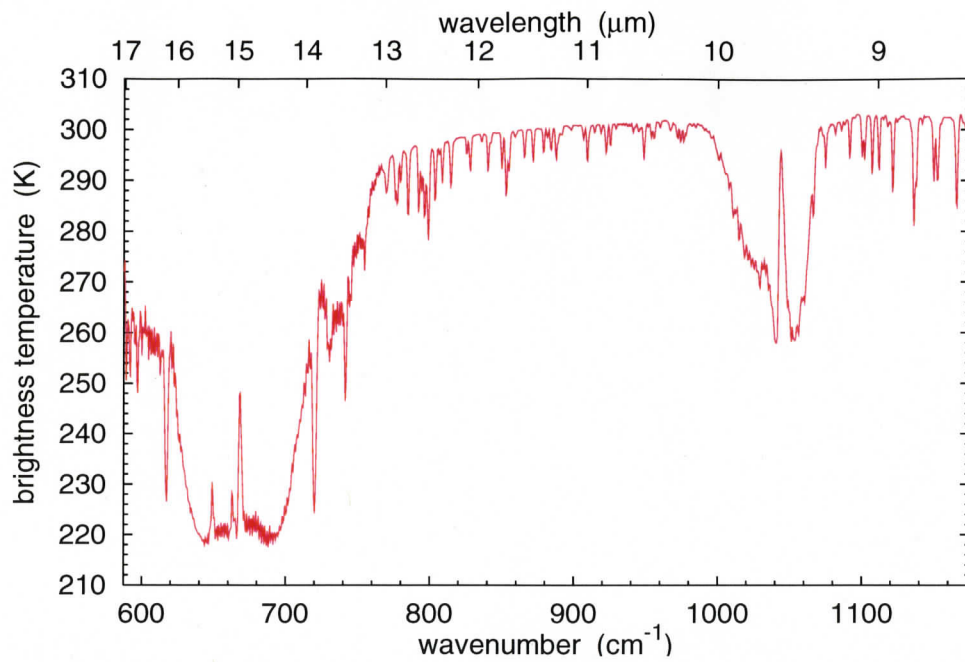


Figure 10: GIFTSFRTE LW brightness temperature spectrum for the clear atmospheric profile of Fig. 9, nadir viewing.

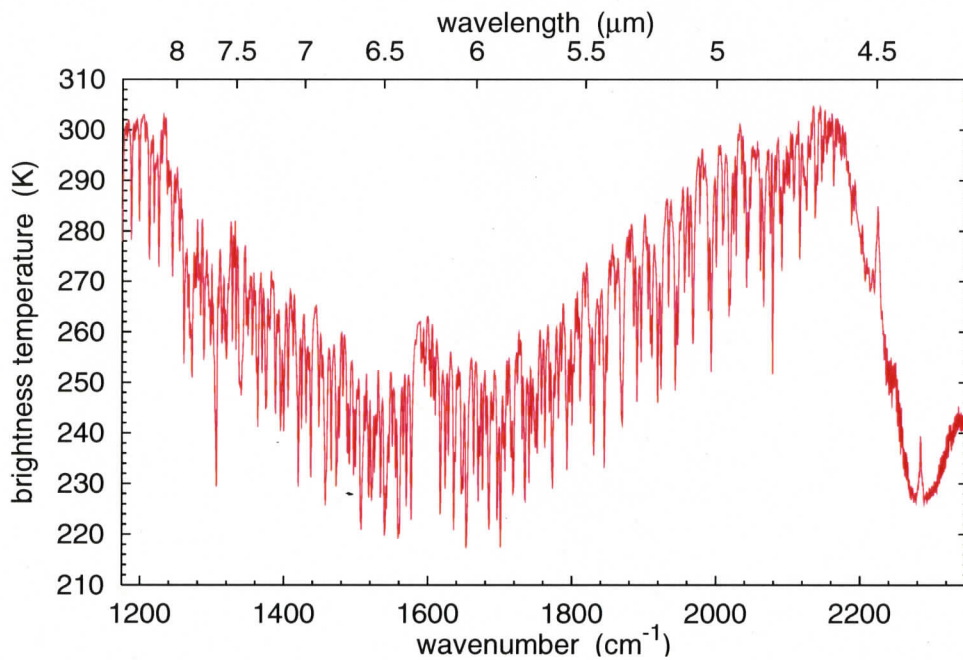


Figure 11: GIFTSFRTE SMW brightness temperature spectrum for the clear atmospheric profile of Fig. 9, nadir viewing.



## 4.2 Effective Layer Temperature

There are a number of techniques for calculating the radiance emergent from an atmospheric layer across which there is a temperature gradient. Since the Planck radiance is a function of temperature, an important consideration is how a single temperature is assigned to an atmospheric layer. Three methods of computing the effective temperature of an atmospheric layer are discussed and compared in this section.

The version of GIFTSEFRTE delivered on 7 January 2003 (along with the Yang:2003 cloud model) has the upwelling radiance contribution from each atmospheric layer computed using the optical depth of the layer and its lower boundary temperature, call this Method 1. A more accurate alternative is to use the layer optical depth and the arithmetic mean Planck radiance,  $\frac{1}{2}[B(T_n) + B(T_f)]$ , where  $T_n$  and  $T_f$  are the temperatures of the near and far boundaries. Call this Method 2. For the example clear atmosphere of Fig. 9, there can be differences in top-of-atmosphere brightness temperatures of up to approximately 1 K between the two methods, as shown in Fig. 12 and Fig. 13.

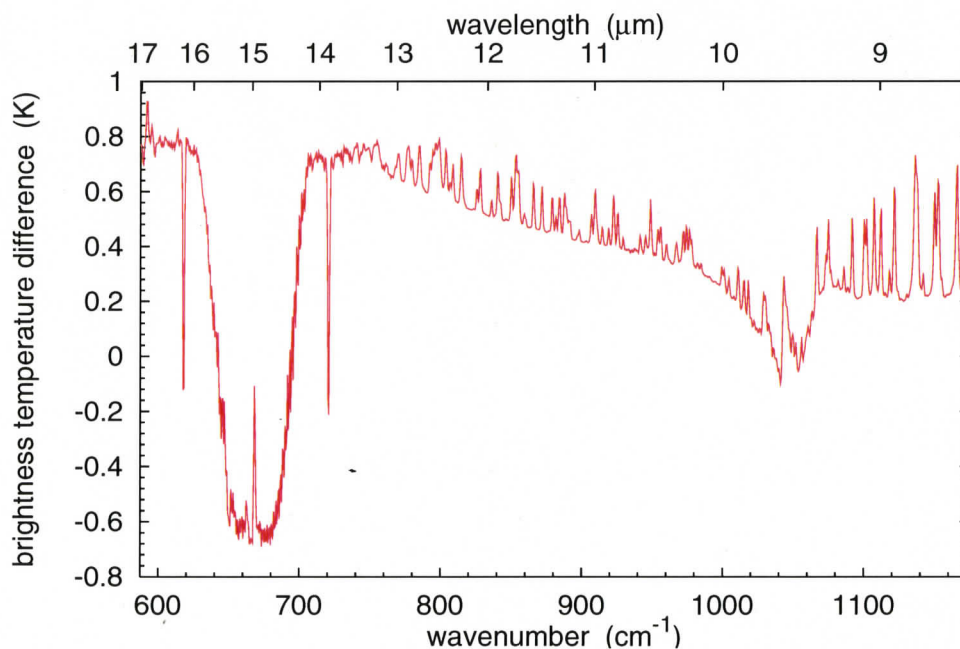


Figure 12: GIFTS LW band Method 1 brightness temperature less Method 2 brightness temperature for atmospheric profile of Fig. 9, nadir viewing.

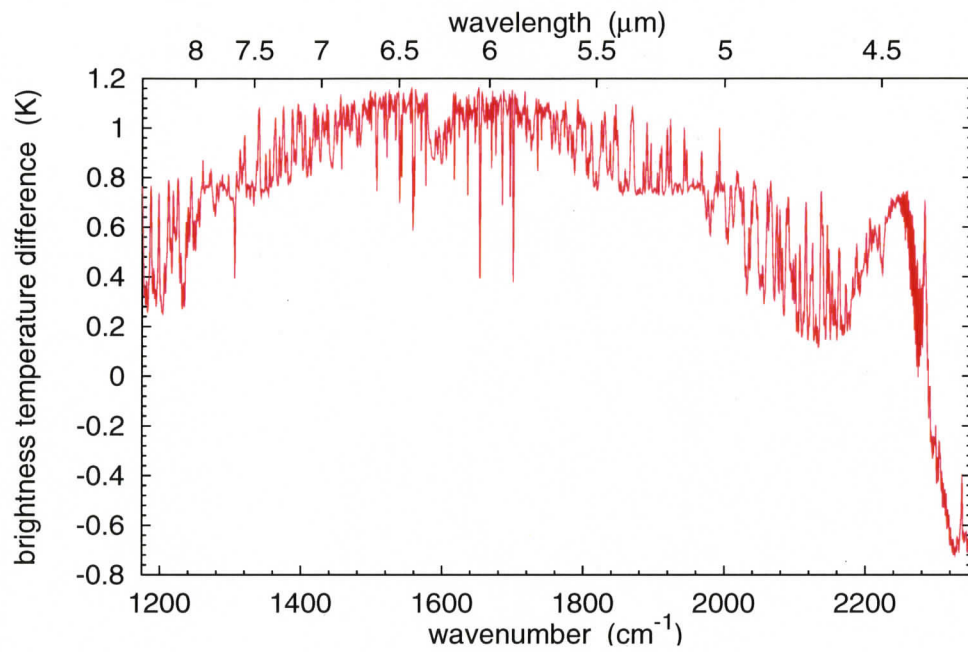


Figure 13: GIFTS SMW band Method 1 brightness temperature less Method 2 brightness temperature for atmospheric profile of Fig. 9, nadir viewing.

Following the suggestion of Dave Turner<sup>2</sup>, a third method (Method 3) was also investigated, in which a two term Padé approximation to the effective Planck function of an atmospheric layer (of optical depth,  $\delta$ ) is given by, [1],

$$B(T_e) = \frac{\frac{1}{2}[B(T_n) + B(T_f)] + (a\delta + b\delta^2) B(T_n)}{1 + a\delta + b\delta^2}, \quad (1)$$

where  $T_e$  is the layer effective temperature. In Eqn. (1)  $a = 0.193$  and  $b = 0.013$ .

The brightness temperature differences between Method 3 and Method 2 for nadir viewing are shown in Fig. 14 and Fig. 15. There is only a small difference in brightness temperature, not more than 0.1 K, depending on channel. Fig. 16 and Fig. 17 show that at an observation zenith of  $60^\circ$  this small difference is increased only a small amount compared to nadir viewing. On the authority of [1], Method 3 is selected as the operational algorithm for GIFTSFRTE.

<sup>2</sup>Pacific Northwest National Laboratories

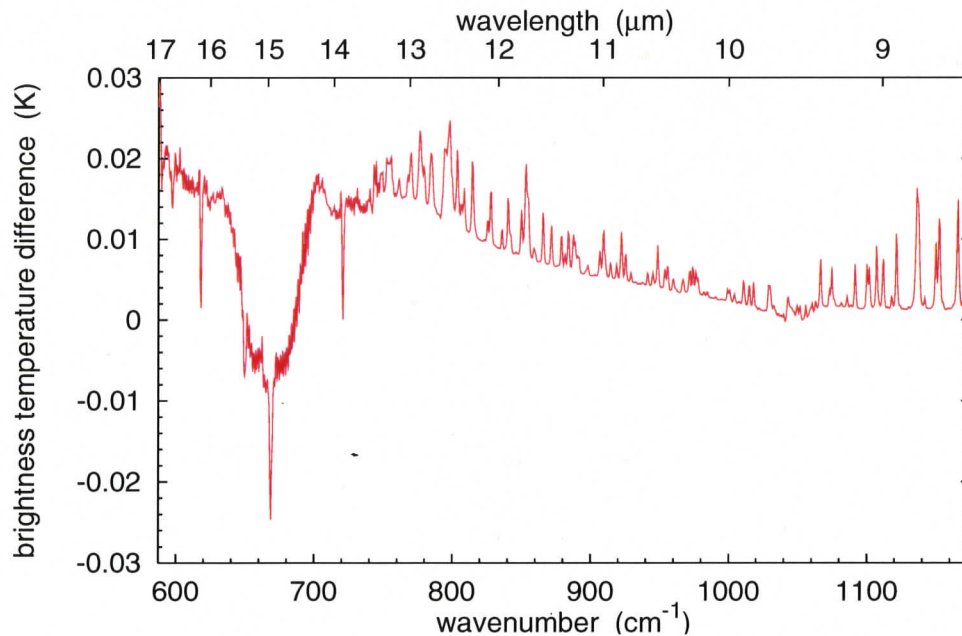


Figure 14: GIFTS LW band Method 3 brightness temperature less Method 2 brightness temperature for atmospheric profile of Fig. 9, nadir viewing.

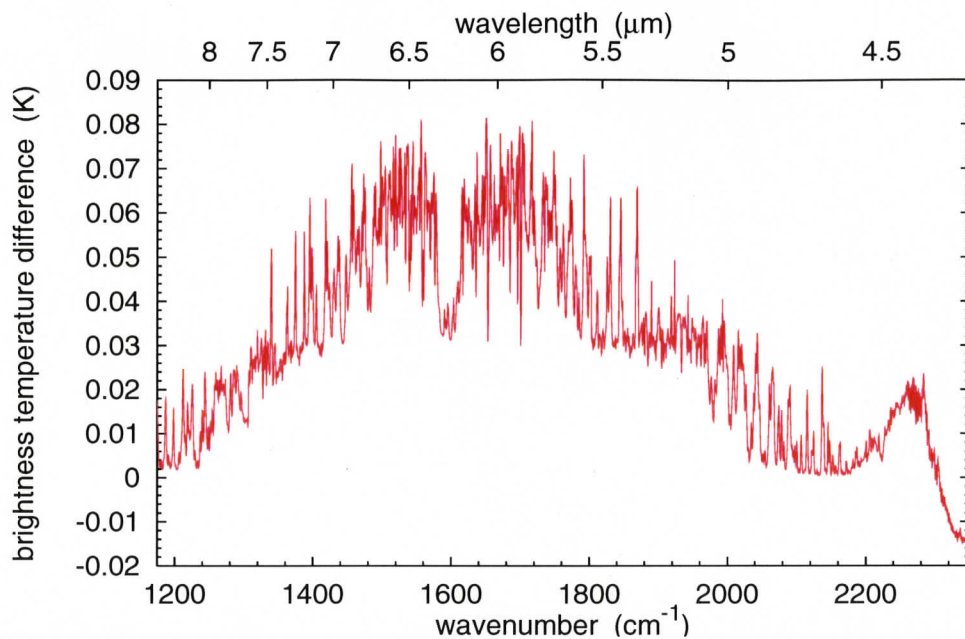


Figure 15: GIFTS SMW band Method 3 brightness temperature less Method 2 brightness temperature for atmospheric profile of Fig. 9, nadir viewing.

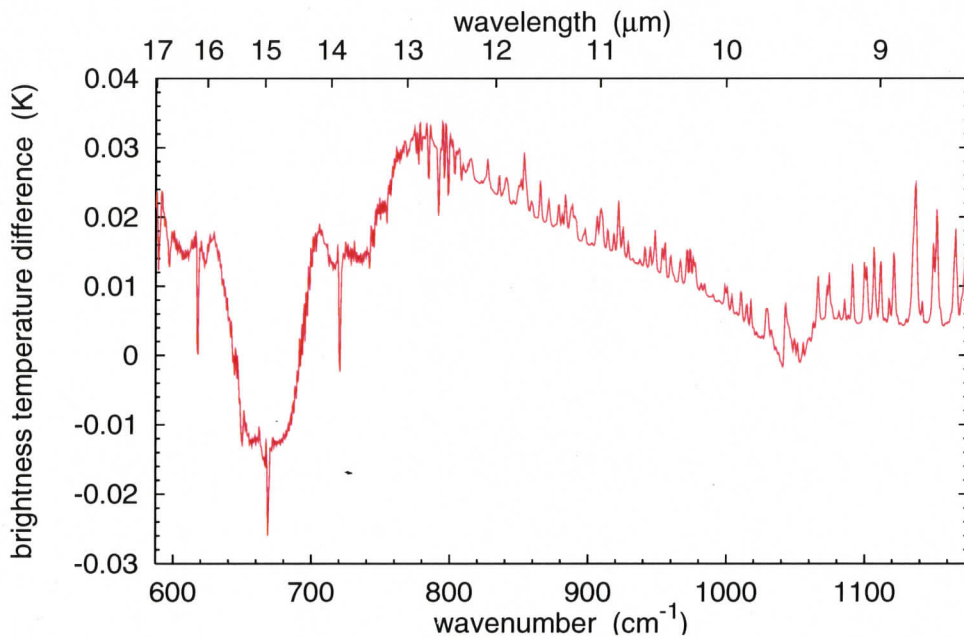


Figure 16: GIFTS LW band Method 3 brightness temperature less Method 2 brightness temperature for atmospheric profile of Fig. 9, 60 deg observation zenith angle.

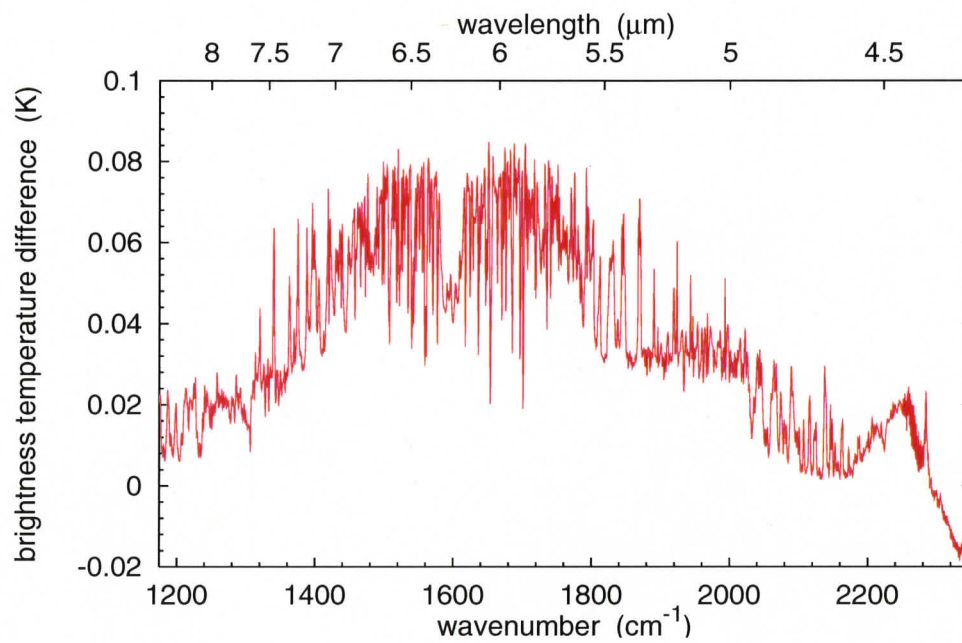


Figure 17: GIFTS SMW band Method 3 brightness temperature less Method 2 brightness temperature for atmospheric profile of Fig. 9, 60 deg observation zenith angle.

### 4.3 Cloud Particle Effective Diameter

The Yang:2003 cloud model in GIFTSFRTE accepts inputs such as the effective diameter of cloud droplets, the cloud phase (liquid or ice), the optical thickness of the cloud at visible wavelengths and the pressure level at the cloud top. The model can accommodate a single cloud layer of either ice crystals or liquid water droplets. The mesoscale model MM5 can deliver the concentrations and effective diameters of five condensate types (two liquid, three ice) at 101 atmospheric levels, i.e. complete profiles. The two liquid condensates are denoted “rain” and “liquid”, and the three ice condensates are “ice”, “snow” and “graupel”.

Fig. 18 shows the atmospheric profile where there is both liquid water and ice cloud present. The pressure levels denoted `pliq` and `pice` are the designated cloud top pressures for liquid water and ice clouds respectively; they are the lowest pressures at which a mass-in-mass mixing ratio of  $1 \times 10^{-6}$  is observed.

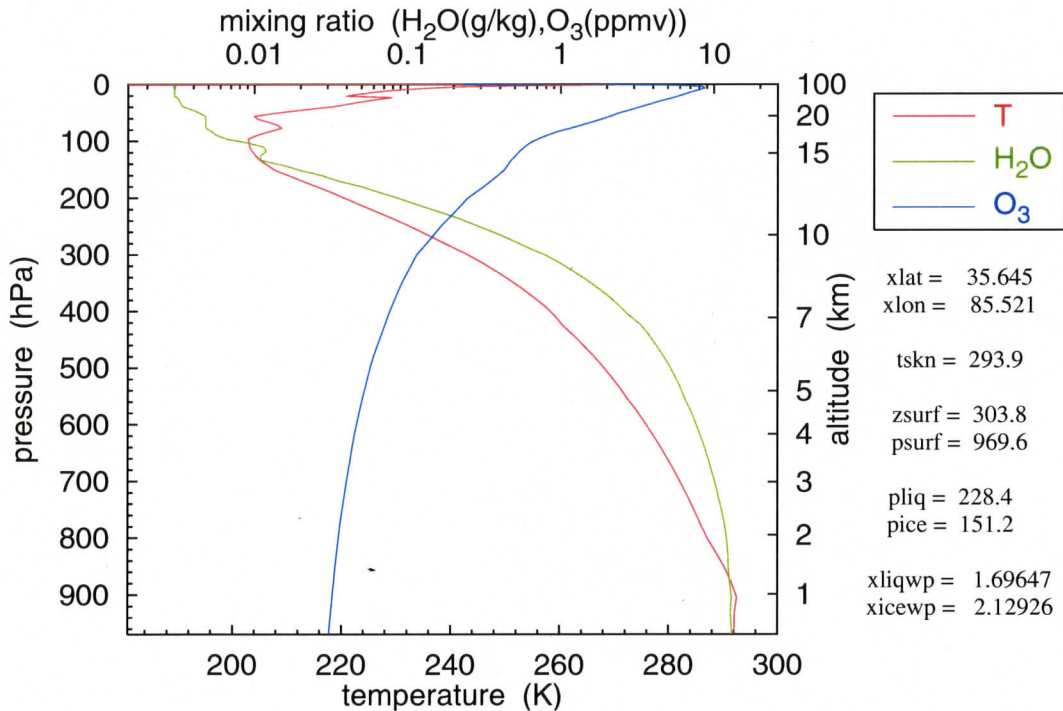


Figure 18: Atmospheric profiles from GIFTS data cube 06122002.2200\_2.3 at (-31,-32). At this location the liquid water path (`xliqwp`) is  $1.6965 \text{ g cm}^{-2}$  and the ice water path (`xicewp`) is  $2.1293 \text{ g cm}^{-2}$ .

Fig. 19 and Fig. 20 show the condensate mixing ratio and effective diameter profiles, respectively.

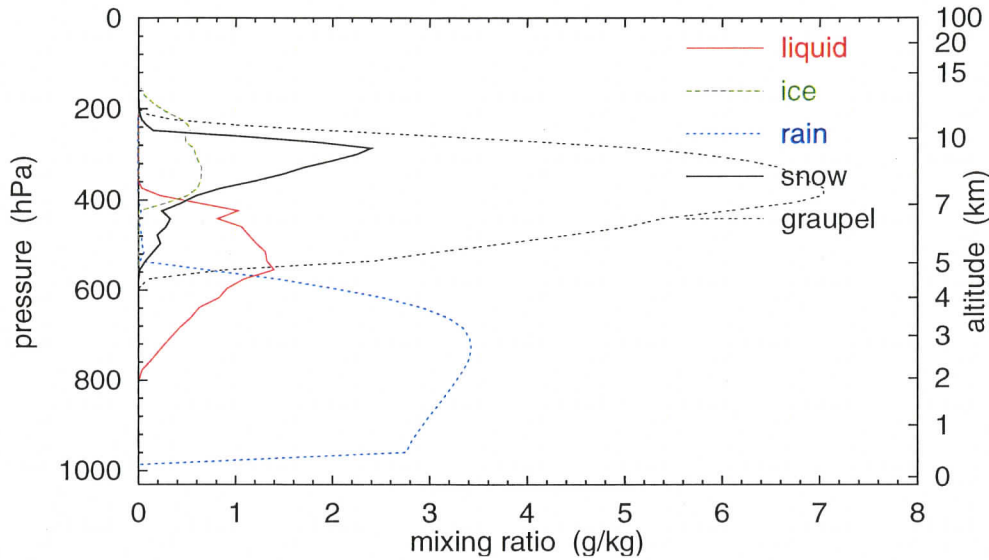


Figure 19: Condensate mixing ratio profile from GIFTS data cube 06122002.2200.2.3 at  $(-31,-32)$ .

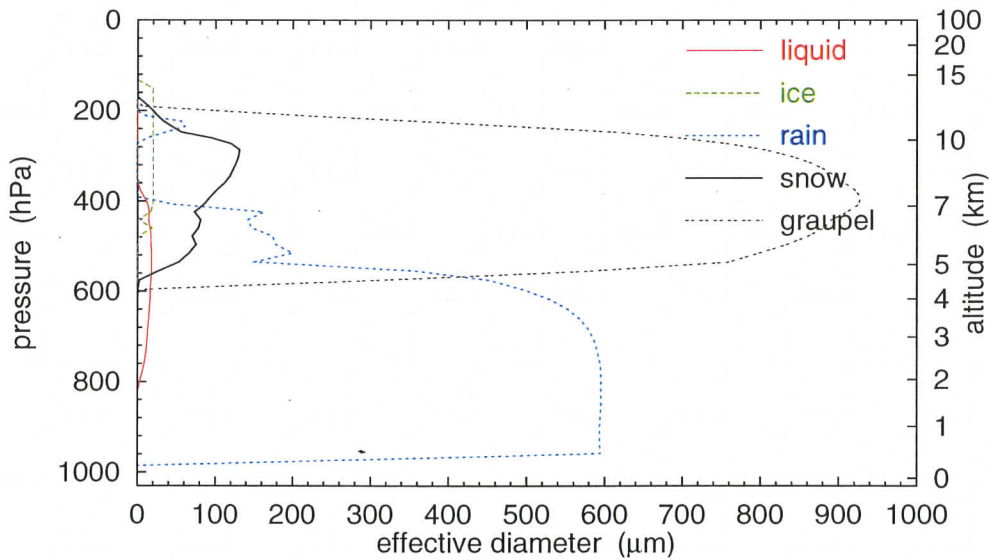


Figure 20: Effective diameter profile from GIFTS data cube 06122002.2200.2.3 at  $(-31,-32)$ . The larger particle sizes are beyond the bounds of applicability of the new cloud model ( $100 \mu\text{m}$  for liquid,  $150 \mu\text{m}$  for ice); this example was selected because it shows all five condensate species.

The condensate profile data must be pre-processed to provide an estimate of the model input parameters of cloud phase, effective diameter and optical depth. The following describes how this is achieved. Section 5 on page 24 assesses the efficacy of this simple approach.

#### 4.3.1 Liquid phase cloud

In the Yang:2003 cloud model, liquid water clouds are assumed to contain spherical droplets with a size distribution based on Deirmendjian's modified gamma distribution [2]. The effective size of liquid cloud droplets,  $D_l$ , is [9],

$$D_l = \frac{\int_0^\infty n(D) D^3 dD}{\int_0^\infty n(D) D^2 dD} \quad , \quad (2)$$

where  $n(D)$  is the number density in the diameter range  $D$  to  $D + dD$ . For an external mixture of "liquid" and "rain" (these are the names given to the two liquid phase condensate size distributions in the MM5 output files),

$$D_l = \frac{\int_0^\infty [n_{rain}(D) + n_{liquid}(D)] D^3 dD}{\int_0^\infty [n_{rain}(D) + n_{liquid}(D)] D^2 dD} \quad . \quad (3)$$

Presently the mixing ratio,  $M$ , and the effective diameter,  $D$ , are available to `gifsftre` from the MM5 condensate profiles. The effective diameter of a mixture of "liquid" and "rain" is computed as,

$$D_l = \frac{M_{rain} + M_{liquid}}{M_{rain}/D_{rain} + M_{liquid}/D_{liquid}} \quad , \quad (4)$$

where  $M$  and  $D$  are interpolated to pressure `pliq`.  $M$ , the mass of condensate per unit mass of moist air, has replaced the condensate volume in the numerator of Eqn. (3) since  $M$  is proportional to condensate volume for air masses at the same atmospheric pressure, i.e. `pliq` or `pice`.

#### 4.3.2 Ice phase cloud

The effective size of hexagonal ice crystals is [4,9],

$$D_i = \frac{3}{2} \times \frac{\text{volume}}{\text{projected area}} = \frac{3\sqrt{3}a^2L}{\sqrt{3}a^2 + 2aL} \quad , \quad (5)$$

where  $a$  is the half-width of cross section and  $L$  is the crystal length (i.e. the longest dimension). The crystal aspect ratio,  $\frac{2a}{L}$ , is related to  $L$  via [9],



$$\frac{2a}{L} = \begin{cases} 1 & , & L \leq 40 \mu m & , \\ \exp[-0.017835(L - 40)] & , & 40 < L \leq 50 \mu m & , \\ \frac{5.916}{\sqrt{L}} & , & L > 50 \mu m & . \end{cases} \quad (6)$$

If the effective diameters of ice phase condensates provided by MM5 are derived from expressions similar to Eqn. (5), it is not possible to recover  $L$  and  $a$  in order to generate an appropriately weighted effective diameter,  $D_i$ . For present purposes,

$$D_i = \frac{M_{ice} + M_{snow} + M_{graupel}}{M_{ice}/D_{ice} + M_{snow}/D_{snow} + M_{graupel}/D_{graupel}} \quad , \quad (7)$$

where  $M$  is the mass mixing ratio and  $D$  the effective diameter from the MM5 condensate profile interpolated to pressure `pice`. “ice”, “snow” and “graupel” are the names given to the three ice phase condensate size distributions in the MM5 output files.

#### 4.4 Selection Rule

Given that the cloud model delivered for GIFTSFRTE can include a single layer cloud of liquid water or ice, but not both, a selection rule must be invoked in the presence of mixed phase cloud and multi-layer clouds. The selection rule is simply that the cloud phase found at the highest altitude (lowest pressure) is the one used for forward model simulations. The optical depth is determined by the column amount of that phase but the effective diameter of particles is drawn from the condensate profile interpolated to the nominated cloud top pressure.

## 5 VERIFICATION

To verify that GIFTSFRTE is operating correctly we employ LBLRTM [1] to generate layer gaseous optical depths. These are merged with the single scattering properties of clouds which are then processed by DISORT [6] to generate simulated radiances at the top of the atmosphere at a specified spectral resolution. LBLDIS is a computer code which facilitates this task [7, 8].

Spectral reduction to GIFTS channel radiances and conversion to brightness temperatures permit comparisons to be made with the GIFTS forward model. Spectral reduction is achieved by the MATLAB code `gifts.m` written by Dave Tobin.

The atmosphere is divided into 101 layers for LBLRTM (and, subsequently, DISORT) calculations. This is to match the layering in GIFTSFRTE which has been hard-coded to the 101 pressure levels defined for retrievals from the Atmospheric Infra-Red Sounder (AIRS). These 101 pressure levels are shown in Table 1. The atmospheric profiles initially used to generate GIFTS radiances (and brightness temperatures) are the LOWTRAN standard atmospheres which are defined only to 1013 hPa. As a consequence AIRS layers 1, 2 and 3 will normally not be used except under conditions of high barometric pressure and over land surfaces that are below sea level.

Brightness temperatures computed in this way are compared to GIFTSFRTE brightness temperatures that are generated for the same atmospheric profile. The cloud model in GIFTSFRTE can simulate only single-layer, single-phase clouds so that, in the first instance, only equivalently idealised profiles are provided to LBLDIS. That is, cloud optical properties are merged into a single AIRS layer in the LBLDIS processing, and reflection and transmission functions for a cloud layer (at the top of the given AIRS layer) in the GIFTSFRTE processing.

The procedure to verify GIFTSFRTE can be summarised as,

1. execute LBLRTM in optical depth mode to generate layer optical depths for a standard clear atmosphere (layers defined using the AIRS pressure levels),
2. define a cloud layer in terms of its altitude, type (i.e. liquid or ice) and moments of its size distribution
3. execute LBLDIS to generate radiances at some specified resolution and spectrally reduce to GIFTS channels radiances, convert to brightness temperatures,
4. construct an atmosphere in the GIFTS data cube format which is equivalent to the standard atmosphere used by LBLRTM/DISORT plus a single cloud layer occupying one of the AIRS layers,
5. execute GIFTSFRTE and compare brightness temperatures computed by the fast model (GIFTSFRTE) and the verification code (LBLRTM/LBLDIS), and

6. repeat items 2 to 5 above for both liquid water and ice clouds.

layer	$P$ (hPa)	OD file	layer	$P$ (hPa)	OD file
101	0.0000 — 0.0050	ODdeft_098			
100	0.0050 — 0.0161	ODdeft_097	50	151.2664 — 160.4959	ODdeft_047
99	0.0161 — 0.0384	ODdeft_096	49	160.4959 — 170.0784	ODdeft_046
98	0.0384 — 0.0769	ODdeft_095	48	170.0784 — 180.0183	ODdeft_045
97	0.0769 — 0.1370	ODdeft_094	47	180.0183 — 190.3203	ODdeft_044
96	0.1370 — 0.2244	ODdeft_093	46	190.3203 — 200.9887	ODdeft_043
95	0.2244 — 0.3454	ODdeft_092	45	200.9887 — 212.0277	ODdeft_042
94	0.3454 — 0.5064	ODdeft_091	44	212.0277 — 223.4415	ODdeft_041
93	0.5064 — 0.7140	ODdeft_090	43	223.4415 — 235.2338	ODdeft_040
92	0.7140 — 0.9753	ODdeft_089	42	235.2338 — 247.4085	ODdeft_039
91	0.9753 — 1.2972	ODdeft_088	41	247.4085 — 259.9691	ODdeft_038
90	1.2972 — 1.6872	ODdeft_087	40	259.9691 — 272.9191	ODdeft_037
89	1.6872 — 2.1526	ODdeft_086	39	272.9191 — 286.2617	ODdeft_036
88	2.1526 — 2.7009	ODdeft_085	38	286.2617 — 300.0000	ODdeft_035
87	2.7009 — 3.3398	ODdeft_084	37	300.0000 — 314.1369	ODdeft_034
86	3.3398 — 4.0770	ODdeft_083	36	314.1369 — 328.6753	ODdeft_033
85	4.0770 — 4.9204	ODdeft_082	35	328.6753 — 343.6176	ODdeft_032
84	4.9204 — 5.8776	ODdeft_081	34	343.6176 — 358.9665	ODdeft_031
83	5.8776 — 6.9567	ODdeft_080	33	358.9665 — 374.7241	ODdeft_030
82	6.9567 — 8.1655	ODdeft_079	32	374.7241 — 390.8926	ODdeft_029
81	8.1655 — 9.5119	ODdeft_078	31	390.8926 — 407.4738	ODdeft_028
80	9.5119 — 11.0038	ODdeft_077	30	407.4738 — 424.4698	ODdeft_027
79	11.0038 — 12.6492	ODdeft_076	29	424.4698 — 441.8819	ODdeft_026
78	12.6492 — 14.4559	ODdeft_075	28	441.8819 — 459.7118	ODdeft_025
77	14.4559 — 16.4318	ODdeft_074	27	459.7118 — 477.9607	ODdeft_024
76	16.4318 — 18.5847	ODdeft_073	26	477.9607 — 496.6298	ODdeft_023
75	18.5847 — 20.9224	ODdeft_072	25	496.6298 — 515.7200	ODdeft_022
74	20.9224 — 23.4526	ODdeft_071	24	515.7200 — 535.2322	ODdeft_021
73	23.4526 — 26.1829	ODdeft_070	23	535.2322 — 555.1669	ODdeft_020
72	26.1829 — 29.1210	ODdeft_069	22	555.1669 — 575.5248	ODdeft_019
71	29.1210 — 32.2744	ODdeft_068	21	575.5248 — 596.3062	ODdeft_018
70	32.2744 — 35.6505	ODdeft_067	20	596.3062 — 617.5112	ODdeft_017
69	35.6505 — 39.2566	ODdeft_066	19	617.5112 — 639.1398	ODdeft_016
68	39.2566 — 43.1001	ODdeft_065	18	639.1398 — 661.1920	ODdeft_015
67	43.1001 — 47.1882	ODdeft_064	17	661.1920 — 683.6673	ODdeft_014
66	47.1882 — 51.5278	ODdeft_063	16	683.6673 — 706.5654	ODdeft_013
65	51.5278 — 56.1260	ODdeft_062	15	706.5654 — 729.8857	ODdeft_012
64	56.1260 — 60.9895	ODdeft_061	14	729.8857 — 753.6275	ODdeft_011
63	60.9895 — 66.1253	ODdeft_060	13	753.6275 — 777.7897	ODdeft_010
62	66.1253 — 71.5398	ODdeft_059	12	777.7897 — 802.3714	ODdeft_009
61	71.5398 — 77.2396	ODdeft_058	11	802.3714 — 827.3713	ODdeft_008
60	77.2396 — 83.2310	ODdeft_057	10	827.3713 — 852.7880	ODdeft_007
59	83.2310 — 89.5204	ODdeft_056	9	852.7880 — 878.6201	ODdeft_006
58	89.5204 — 96.1138	ODdeft_055	8	878.6201 — 904.8659	ODdeft_005
57	96.1138 — 103.0172	ODdeft_054	7	904.8659 — 931.5236	ODdeft_004
56	103.0172 — 110.2366	ODdeft_053	6	931.5236 — 958.5911	ODdeft_003
55	110.2366 — 117.7775	ODdeft_052	5	958.5911 — 986.0666	ODdeft_002
54	117.7775 — 125.6456	ODdeft_051	4	986.0666 — 1013.948	ODdeft_001 <sup>†</sup>
53	125.6456 — 133.8462	ODdeft_050	3	1013.948 — 1042.232	N/A
52	133.8462 — 142.3848	ODdeft_049	2	1042.232 — 1070.917	N/A
51	142.3848 — 151.2664	ODdeft_048	1	1070.917 — 1100.000	N/A

Table 1: 101 AIRS atmospheric layers. The “OD file” is the filename of gaseous spectral optical depths computed by LBLRTM. The atmospheric profiles initially used to generate GIFTS radiances (and brightness temperatures) are the LOW-TRAN standard atmospheres which are defined only to 1013 hPa.

<sup>†</sup>NOTE: The lower boundary of ODdeft\_001 is at pressure level 1013 hPa *not* the AIRS pressure level 1013.948 hPa.

## 5.1 Clear Soundings

The procedure outlined above for verification of GIFTSFRTTE first requires that GIFTS channel radiances from LBLRTM and from LBLDIS be in agreement for clear atmospheres. The method adopted to test this is,

1. execute LBLRTM in radiance mode to simulate nadir observation of a standard atmosphere,
2. extract the monochromatic radiances (using `fsc2asc` from Paul van Delst),
3. spectrally reduce to generate GIFTS channel radiances (using `gifts.m` from Dave Tobin; `aflag=1`: KaiserBessel # 6 apodization),
4. execute LBLRTM in optical depth mode for the same standard atmosphere,
5. execute LBLDIS (using the layer optical depths generated by LBLRTM) to generate radiances at some specified resolution<sup>3</sup> and spectrally reduce to generate GIFTS channel radiances (via `gifts.m`), and
6. convert GIFTS channel radiances to brightness temperatures and compare results.

Fig. 21 shows the brightness temperature from GIFTS LW band radiances simulated with LBLRTM for the US 1976 standard atmosphere. Also shown in Fig. 21 are the differences between this result and the results obtained using LBLDIS for different spectral resolutions. The comparisons show that, for LBLDIS executed with spectral step size  $0.02 \text{ cm}^{-1}$ , discrepancies of as large as 4 K occur in the  $\text{CO}_2$  band centred at  $15 \mu\text{m}$  and discrepancies of about 2 K occur in the  $9.6 \mu\text{m}$   $\text{O}_3$  band. Reducing the spectral step size to  $0.01 \text{ cm}^{-1}$ , which is only possible with the modified LBLDIS code, can reduce discrepancies to about 3 K in the  $\text{CO}_2$  band and about 1 K in the  $\text{O}_3$  band. In the window region between these major absorption features, LBLDIS under-estimates the brightness temperature by about 1 K. Reducing the spectral step size further to  $0.001 \text{ cm}^{-1}$  assists in removing brightness temperature difference fluctuations in the  $\text{O}_3$  band but the magnitude of the discrepancies in the  $\text{CO}_2$  band and the window region remain largely unchanged.

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<sup>3</sup>A minor source code modification was made to LBLDIS to enable the wavenumber bandpass to be reduced below approximately  $0.2 \text{ cm}^{-1}$ . This involved changing layer bandpass transmittances,  $\bar{\tau}$ , from  $\bar{\tau} = \frac{1}{N} \sum_{i=1}^N \tau_i$  to  $\bar{\tau} = \int_a^b \tau(x) dx / (b - a)$  where  $\tau(x)$  is an interpolating function.  $a$  and  $b$  are the wavenumber limits of a boxcar bandpass,  $\tau_i$  are monochromatic transmittances for a given layer taken from LBLRTM monochromatic layer optical depth files.

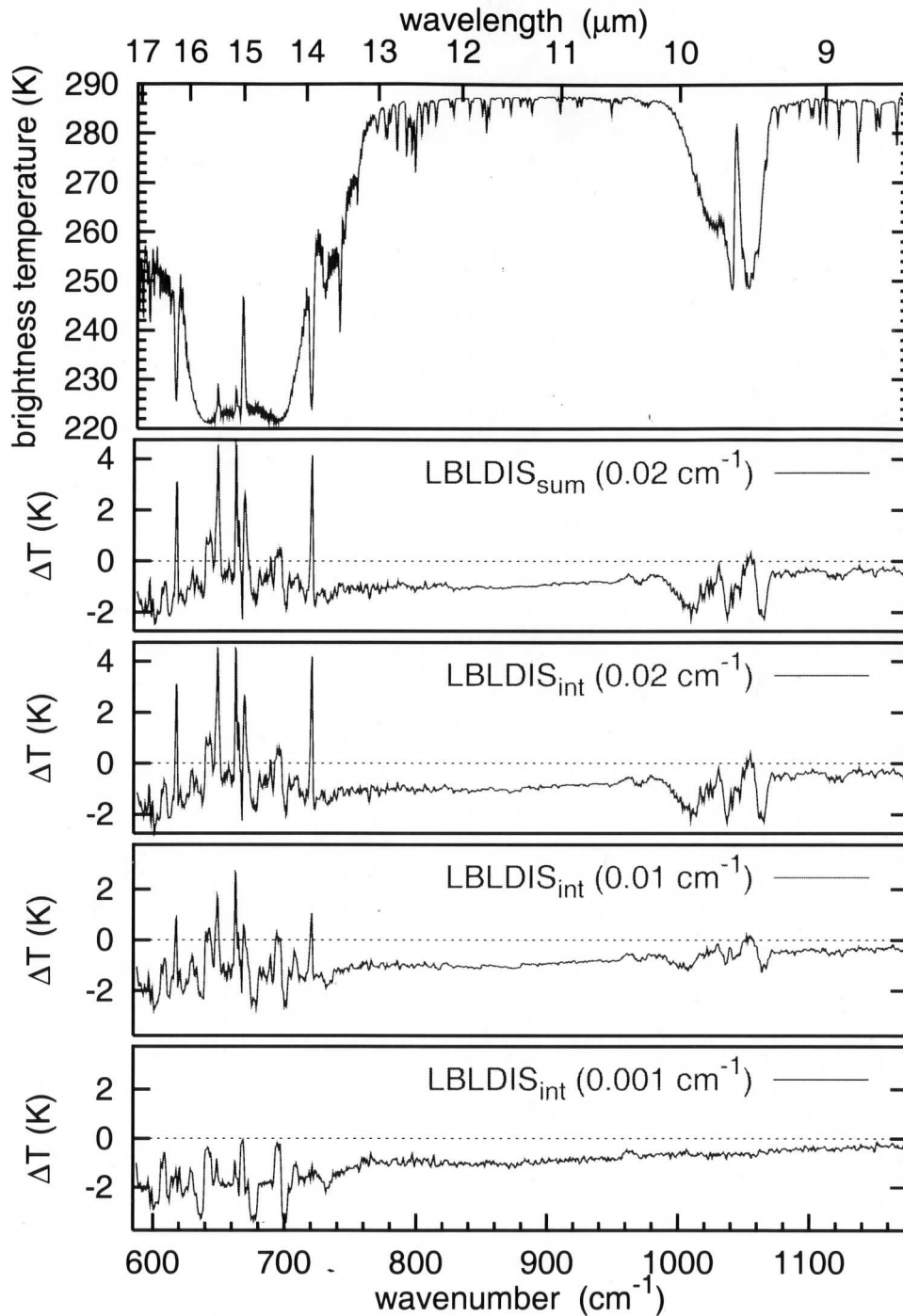


Figure 21: LBLRTM vs DISORT for clear US 1976 standard atmosphere. The upper panel shows the top-of-atmosphere brightness temperature computed by LBLRTM and spectrally reduced to GIFTS LW channels. The panels beneath this show the brightness temperature difference arrived at by subtracting the upper pane result from LBLDIS computations for the same clear atmosphere. Subscript “int” indicates the version of LBLDIS modified according to the footnote of the previous page.

It is thought that the reason for the discrepancies exhibited in Fig. 21 are in the formulation of effective layer optical depth, required by DISORT over a finite width bandpass. The implicit assumption is that the illumination over a narrow bandpass is spectrally flat, an approximation which, for any given bandpass width, becomes increasingly poor with increasing altitude.

In the profile data for which GIFTS radiance data cubes are synthesised there are no clouds with tops above 150 hPa. Consequently, to reduce the discrepancies noted above, LBLDIS is run below the maximum cloud top altitude and LBLRTM separately computes the radiances and transmittances above the maximum cloud top altitude. The LBLDIS and LBLRTM components are then joined in the following manner,

$$L'_i = \text{TRP}[\nu_i^*, M, \nu, L] \times \tau_i^* + L_i^* \quad , \quad (8)$$

where  $L_i^*$  and  $\tau_i^*$  are the monochromatic radiances and transmittances from LBLRTM at wavenumber  $\nu_i^*$ .  $L$  and  $\nu$  are the narrow bandpass radiances and transmittances from LBLDIS and TRP is a linear interpolation function [5]. By way of explanation, if  $X$  and  $Y$  ( $N$  element vectors,  $X$  monotonic) and if  $x$  is a scalar within the span of  $X$ , the statement  $y = \text{TRP}[x, N, X, Y]$  assigns to  $y$  the value  $Y_i + \frac{x - X_i}{X_{i+1} - X_i}(Y_{i+1} - Y_i)$ , where  $x$  falls between the elements  $X_i$  and  $X_{i+1}$ . Spectral reduction of  $L'$  to GIFTS channel radiances (via `gifts.m`) provides the simulated top-of-atmosphere radiances. In the implementation of Eqn. (8) we computed  $L$  from the surface (1013 hPa) to 103.017 hPa, the latter being coincident with an AIRS pressure level and above the level of clouds.  $L_i^*$  and  $\tau_i^*$  are computed from 103.017 hPa to space with the lower boundary characterized by unit emissivity and absolute zero temperature.

Separately computing radiances and transmittances above 103.017 hPa reduces the discrepancy between LBLRTM top-of-atmosphere radiances and those computed with DISORT up to the top-of-atmosphere. Fig. 22 shows differences for the GIFTS LW band and Fig. 23 the differences for the GIFTS SMW band. Discrepancies are typically no greater than 0.1 K except in the 9.6  $\mu\text{m}$  ozone band where discrepancies of 0.15 K can be found.

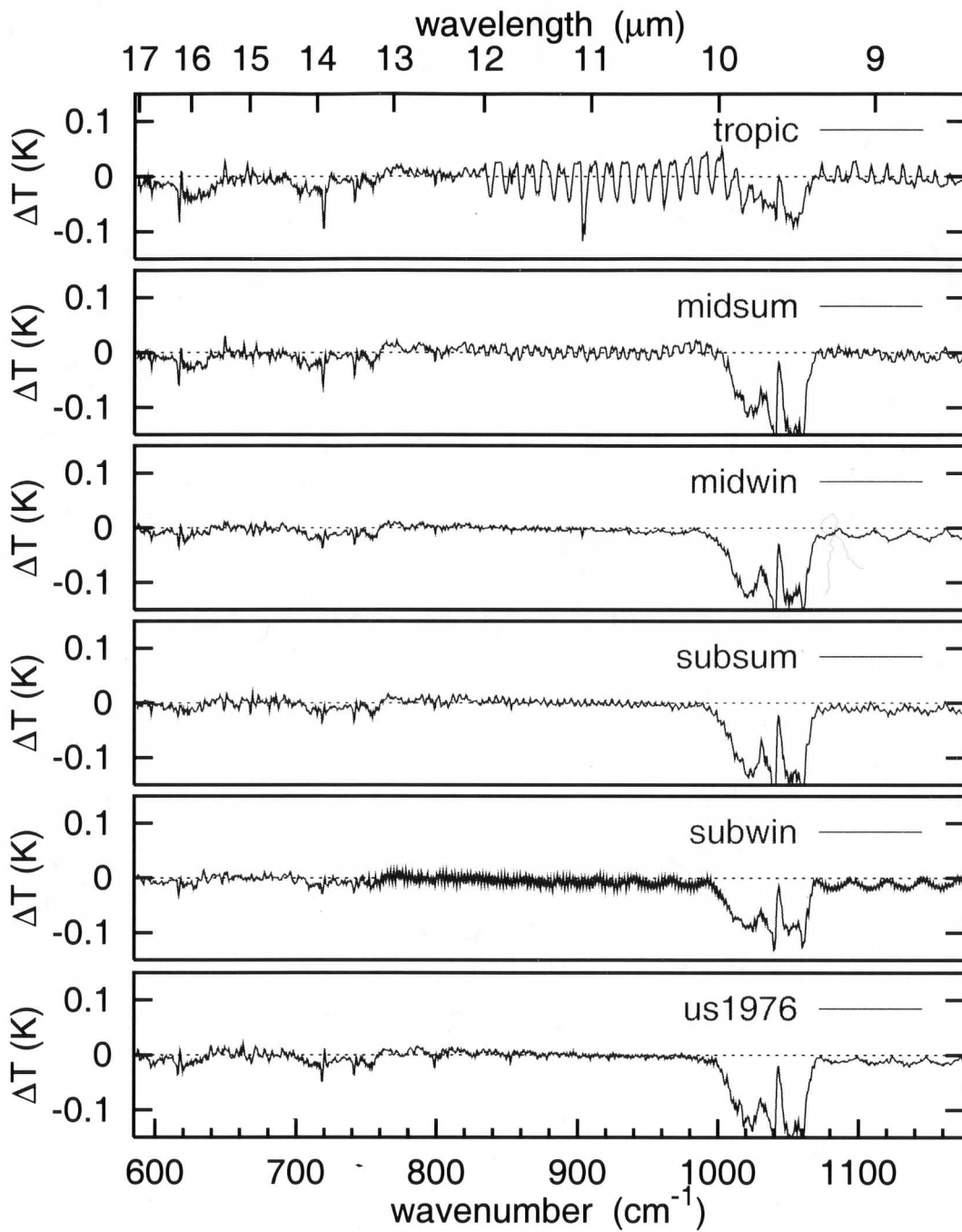


Figure 22: LW brightness temperature differences between DISORT and LBLRTM for the six LOWTRAN standard atmospheres. Discrepancies are typically much less than 0.1 K except in the 9.6  $\mu\text{m}$  ozone band where they may rise to 0.15 K.



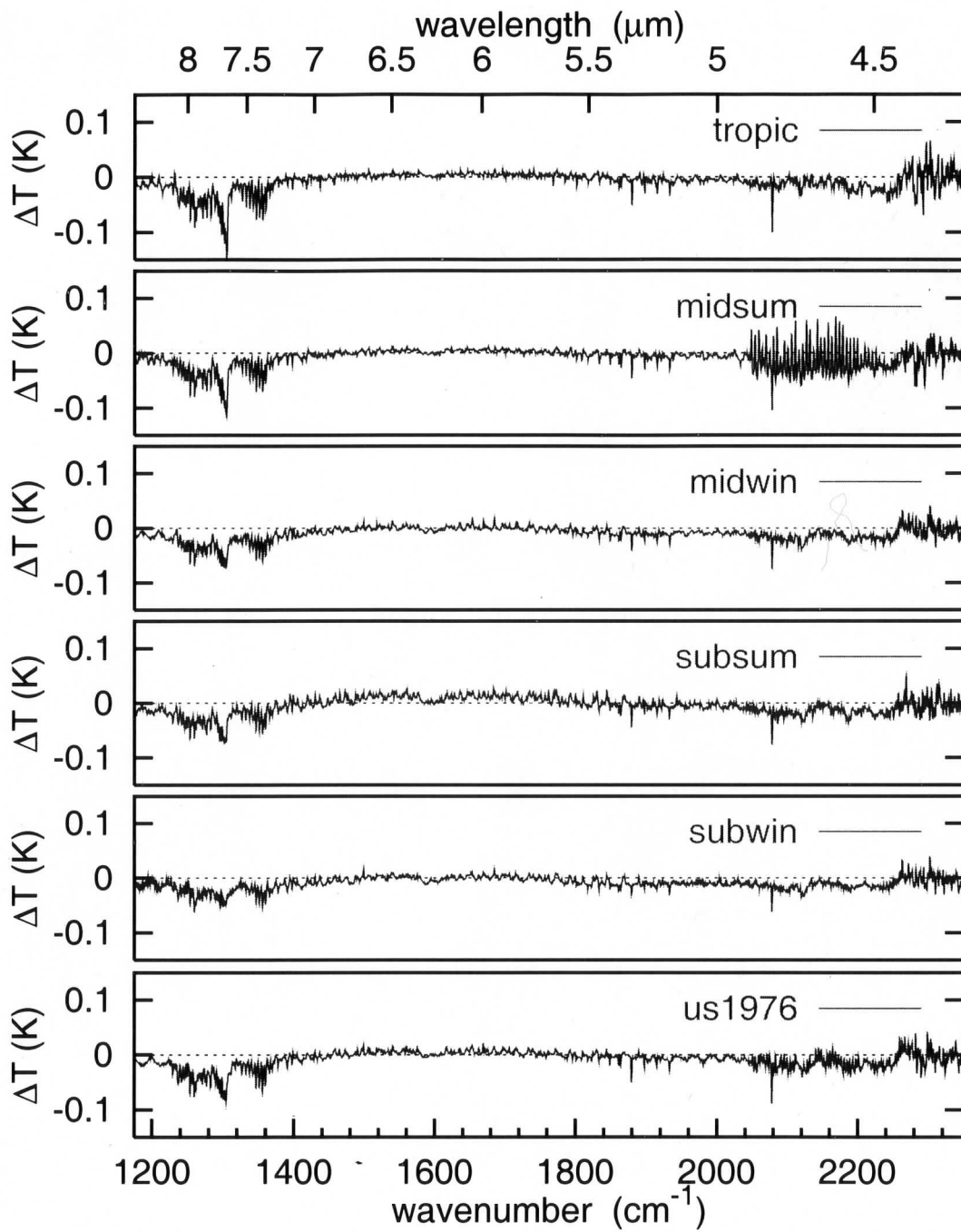


Figure 23: SMW brightness temperature differences between DISORT and LBLRTM for the six LOWTRAN standard atmospheres. Discrepancies are typically of the order 0.05 K and do not rise above 0.15 K in this band.

## 5.2 Cloudy Soundings

Verification of the accuracy of cloudy atmosphere performance consists of comparing top-of-atmosphere brightness temperatures computed by GIFTSFRTE and by LBLDIS for a test set of idealised cloudy atmospheres. For the purposes of verification, the LBLDIS results (obtained in the manner described in Section 5.1) are denoted "TRUTH" and the equivalent GIFTSFRTE results are denoted "FAST".

The atmospheres are idealized in the sense that for TRUTH, cloud droplets are confined to a single AIRS layer and are described by a mono-modal size distribution of variable mode radius but fixed width parameter. For FAST, the cloud layer is defined at the pressure level at the top of the TRUTH layer.

Comparisons are made for liquid and ice clouds, each for four effective diameters, six optical depths and three altitudes. In each case the clear atmosphere profile, to which cloud properties are added, is the US 1976 standard atmosphere.

When the effective diameter of liquid or ice particles is stipulated, rather than computed from a profile of condensate effective sizes and abundances, the size can be entered in the GIFTSFRTE input parameter file. An example input file for GIFTSFRTE is given in Appendix C and for LBLDIS in Appendix D.

### 5.2.1 Liquid phase cloud

For liquid cloud, cloud top altitudes of approximately 1, 2, and 3 km (more precisely, 1.187, 2.176 and 3.199 km to coincide with AIRS pressure levels at 878.62, 777.79 and 683.67 hPa, respectively) were nominated. The effective diameters of liquid droplets, denoted  $D_l$ , chosen for comparison are 2, 10, 20 and 40  $\mu\text{m}$ . Optical depths are 0.1, 0.5, 1, 2, 3, and 5, defined at 10  $\mu\text{m}$ . The Mie extinction efficiencies for liquid droplets illuminated by 10  $\mu\text{m}$  radiation are shown in Table 2.

LBLDIS permits the optical depth to be specified at 10  $\mu\text{m}$  (or at other infrared wavelengths) for a given  $D_l$  but this information is provided to GIFTSFRTE through  $D_l$  and `xliqwp`. As a consequence, to provide GIFTSFRTE with cloudy profiles equivalent to the profile processed by LBLDIS, the liquid water path is computed as,

$$\text{xliqwp} = \frac{D_l \rho_l}{\langle Q_e \rangle} \times \delta_l \quad . \quad (9)$$

Internally, the visible wavelength optical depth  $\delta_l(vis)$ , which is one dimension of the four-dimensional cloud reflection and transmission look-up tables, is computed as,

$D_l$ ( $\mu\text{m}$ )	$\langle Q_e \rangle$
2	0.14368
10	1.13316
20	2.30230
40	2.56237

Table 2: Mie extinction efficiencies  $\langle Q_e \rangle$  for different effective diameters of liquid droplets.  $\langle Q_e \rangle$  is the extinction efficiency as a wavelength of  $10 \mu\text{m}$ ; the angle brackets indicate that this is for an ensemble of particles with sizes distributed about the stipulated effective diameter  $D_l$ .

$$\delta_l(vis) = 2 \times \frac{x_{liqwp}}{D_l \rho_l} \quad , \quad (10)$$

on the assumption that  $\langle Q_e \rangle \approx 2$  at visible wavelengths for particles in the size range of interest.

The brightness temperature comparisons, FAST less TRUTH, in terms of RMS error and average deviation, are shown in Fig. 24 (LW band) and Fig. 25 (SMW band). Significant discrepancies are still present, and these are at their greatest for droplets with  $D_l = 2 \mu\text{m}$  in the vicinity of unit optical depth. At each of the altitudes, 1, 2, and 3 km, there is a temperature differential of approximately 1.5 K across the layer that, in TRUTH simulations, represents the cloud layer (but which in FAST simulations is not present since reflection and transmission properties are assigned to an infinitesimally thin layer).

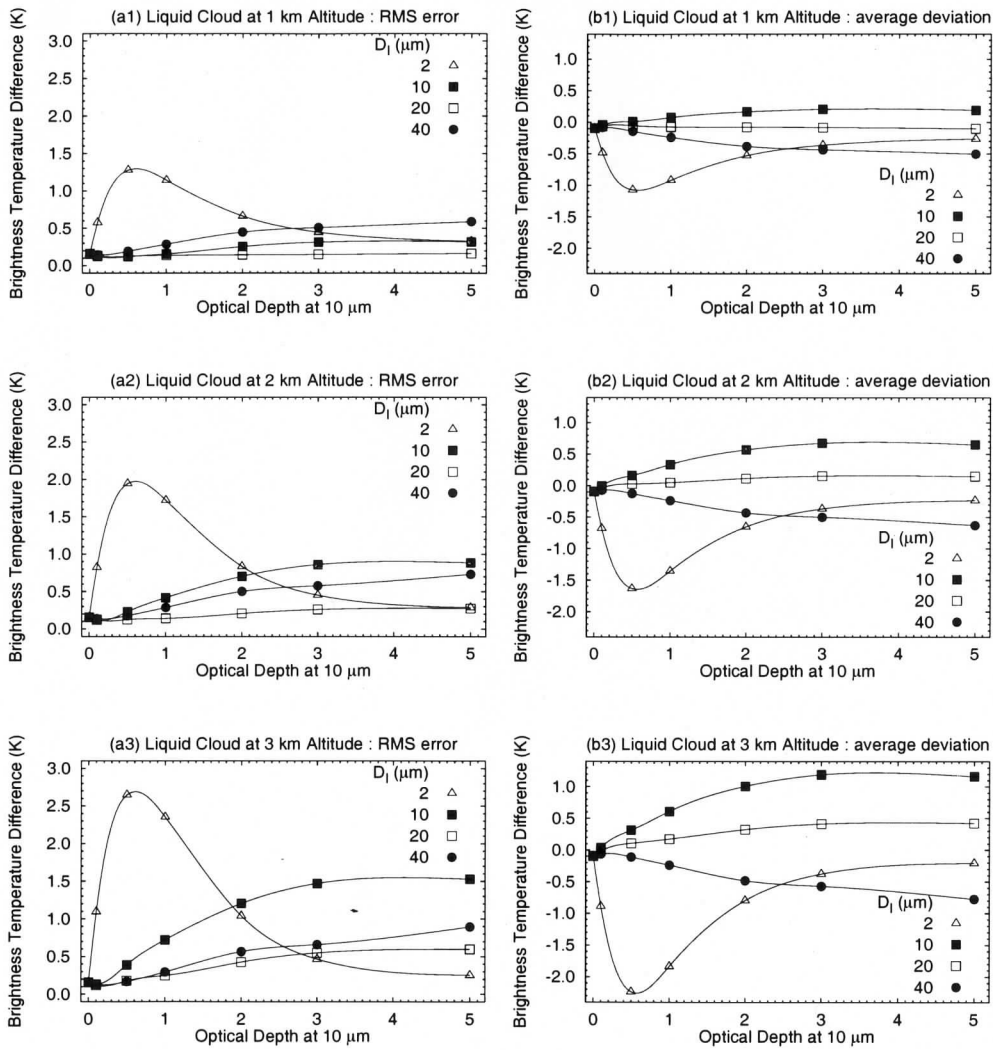


Figure 24: RMS and average deviations (FAST minus TRUTH) for liquid clouds (GIFTS LW).

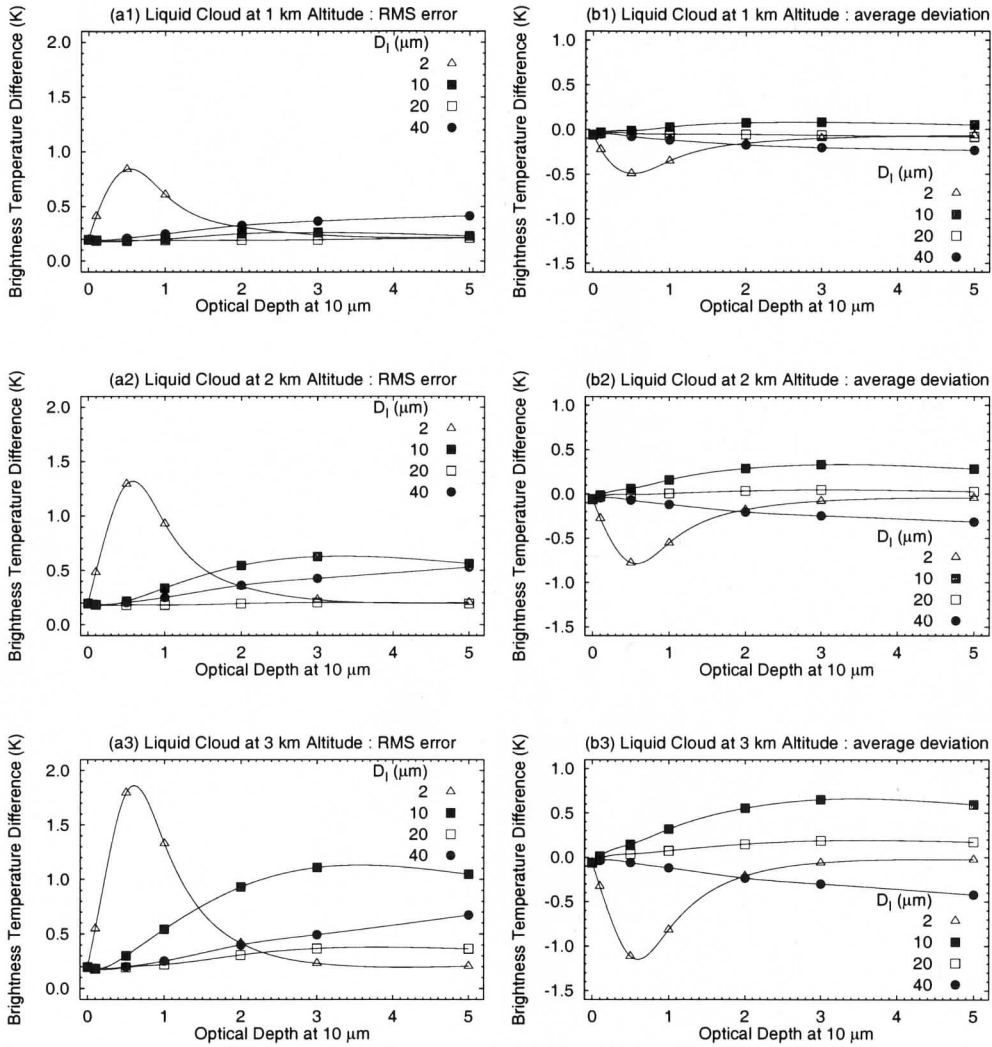


Figure 25: RMS and average deviations (FAST minus TRUTH) for liquid clouds (GIFTS SMW).

$D_i$ ( $\mu\text{m}$ )	$\langle Q_e \rangle$
10	1.06671
20	1.98813
40	2.17719
100	2.06158

Table 3: Extinction efficiencies  $\langle Q_e \rangle$  for different effective diameters of hexagonal ice crystals.

### 5.2.2 Ice phase cloud

Following the procedure for liquid phase cloud, detailed in Section 5.2.1, ice phase cloud top altitudes of approximately 5, 10, and 15 km were nominated (more precisely, 5.073, 10.125 and 15.177 km to coincide with AIRS pressure levels at 535.156, 259.893 and 117.766 hPa, respectively). The effective diameters of liquid droplets chosen for comparison are 10, 20, 40 and 100  $\mu\text{m}$ . Optical depths are 0.1, 0.5, 1, 2, 3, and 5, defined at 10  $\mu\text{m}$ . The extinction efficiencies for hexagonal ice crystals illuminated by 10  $\mu\text{m}$  radiation are shown in Table 3.

LBLDIS permits the optical depth to be specified at 10  $\mu\text{m}$  (or at other infrared wavelengths) for a given  $D_i$  but this information is provided to GIFTSFRTE through  $D_i$  and `xliqwp`. As a consequence, to provide GIFTSFRTE with cloudy profiles equivalent to the profile processed by LBLDIS, the ice water path `xicewp` is computed as,

$$\text{xicewp} = \frac{2}{3} \frac{D_i \rho_l}{\langle Q_e \rangle} \times \delta_i \quad (11)$$

The factor  $\frac{2}{3}$  arises out of consideration of the assymetry of the ice crystal habit [9].

Internally, the visible wavelength optical depth  $\delta_i(vis)$  is computed as,

$$\delta_i(vis) = 3 \times \frac{\text{xicewp}}{D_i \rho_i} \quad (12)$$

on the assumption that  $\langle Q_e \rangle \approx 2$  at visible wavelengths for particles in the size range of interest. `xicewp` is the ice water path in  $\text{g}/\text{m}^2$ ,  $D_i$  is the ice particle effective diameter in  $\mu\text{m}$  and  $\rho_i$  is the density of ice in  $\text{g}/\text{cm}^3$ . We have used  $\rho_i = 0.917 \text{ g}/\text{cm}^3$ .

The brightness temperature comparisons, FAST less TRUTH, in terms of RMS error and average deviation, are shown in Fig. 26 (LW band) and Fig. 27 (SMW band). Discrepancies are still present, which, like the liquid cloud case, are at their greatest for particles at the small end of the range, specifically  $D_i = 10 \mu\text{m}$ , and in the vicinity of two optical depths. At both 5 and 10 km altitudes, there is a temperature differential of approximately 2 K across the layer that, in TRUTH simulations, represents the cloud layer (but which in FAST simulations is, again, not present). However, at 15 km the US 1976 atmosphere is isothermal.

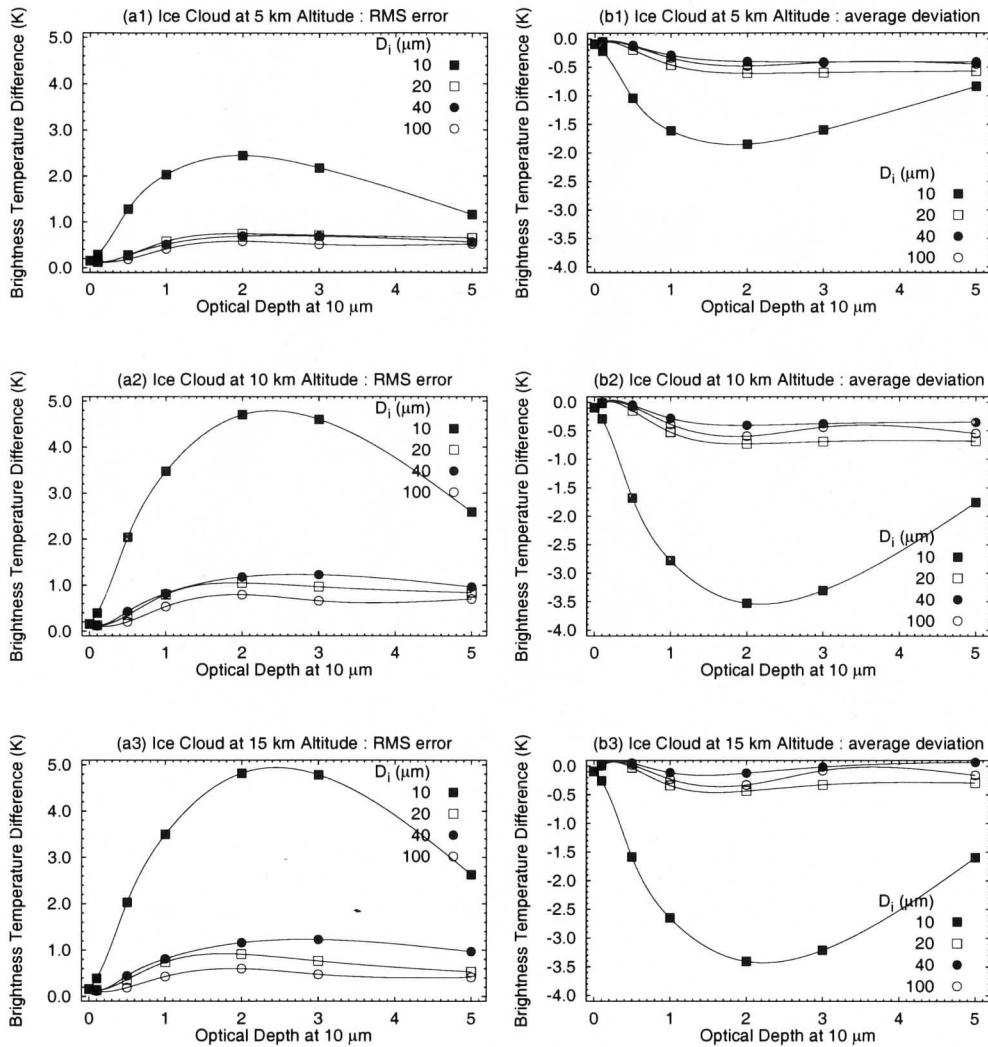


Figure 26: RMS and average deviations (FAST minus TRUTH) for ice clouds (GIFTS LW).

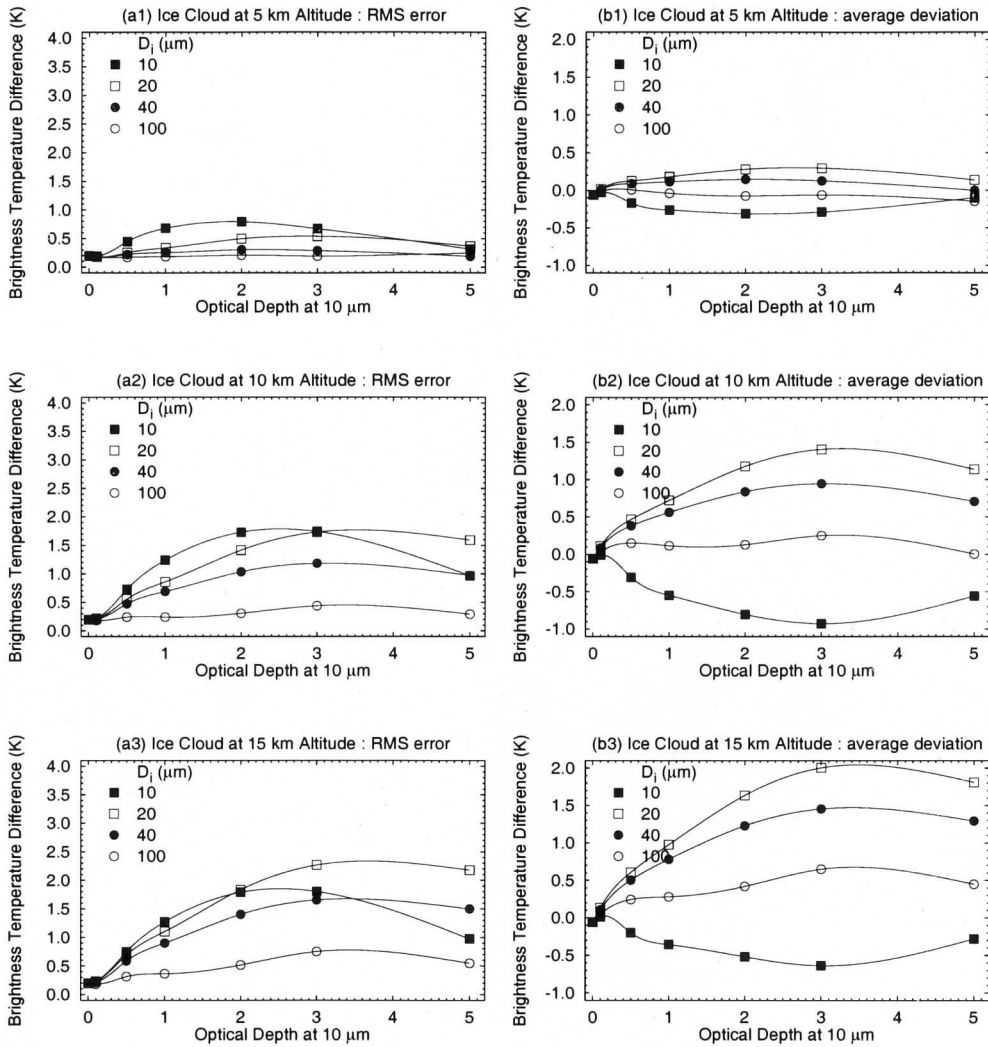


Figure 27: RMS and average deviations (FAST minus TRUTH) for ice clouds (GIFTS SMW).



## 6 CONCLUSION

A new liquid cloud and ice cloud model has been delivered to SSEC for use with the existing GIFTS fast radiative transfer model (GIFTSFRTE). The verification of this code against the more rigorously tested LBLRTM and DISORT shows that discrepancies on the scale of a few degrees Kelvin still exist which cannot be accounted for by the finite geometric thickness of the cloud layer in LBLRTM/DISORT.

There also remain some significant issues with regard to forward modeling of cloudy GIFTS radiances in the presence of mixed phase and multi-level cloud. The cloud model, as delivered, does not seek to account for the impact of cloud below the highest cloud bank. A significant mass of water cloud in the lower troposphere beneath thin cirrus has a spectral emission quite different from thin cirrus alone, the latter being the case that, as presently configured, GIFTSFRTE would simulate.

We are confident, however, that relative changes in GIFTS brightness temperature spectra are more representative of cloud phase and effective diameter than the Old cloud model incorporated into previous versions of GIFTSFRTE.

### Acknowledgements

Thanks to Dave Turner for access to, and advice on, LBLDIS. Thanks also to Brian Osborne for helpful comments on the presentation of this document.

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## A TAPE5 FILE FOR LNFL

Below is the TAPE5 file used to generate a line file using LNFL. The first 7 molecular species in the HITRAN database are used, namely H<sub>2</sub>O, CO<sub>2</sub>, O<sub>3</sub>, N<sub>2</sub>O, CO, CH<sub>4</sub>, and O<sub>2</sub>.

```
aer_hitran_2000_updat_01; 0-10000cm-1; 7 mol; No strength rejection
0.000 10000.000
1111111000000000000000000000000000000000
```

## B TAPE5 FILE FOR LBLRTM

Below is an example TAPE5 file for LBLRTM. In this example, radiances and transmittances are generated for a nadir view of the US1976 standard atmosphere. These values are stored in the binary output file named TAPE12. The surface emissivity is set to unity and the surface temperature is set to 288.2 K which is the boundary layer air temperature. To generate optical depth files, MG on line 4 is set to 1. Otherwise the input file can be left unchanged. The pressure levels stipulated are the AIRS atmospheric pressure levels that are used in the GIFTS forward model. The CO<sub>2</sub> mixing ratio is set to 380 ppmv and freons have not been included.

```

1           2           3           4           5           6           7           8
123456789 123456789 123456789 123456789 123456789 123456789 123456789 123456789
$ Standard atmosphere TAPE5 created by jplot._writetape5
HI=1 F4=1 CN=1 AE=0 EM=1 SC=0 FI=0 PL=0 TS=0 AM=1 MG=0 LA=0 OD=0 XS=0 00 00
500.000 2500.000                0.0002 0.001
288.200 1.000          0.000      0.000      0.000      0.000      0.000
6 2 -99 1 1 7 1 380.000
5.000E-05 1013.000 180.000
1013.000 986.067 958.591 931.524 904.866 878.620 852.788 827.371
802.371 777.790 753.628 729.886 706.565 683.667 661.192 639.140
617.511 596.306 575.525 555.167 535.232 515.720 496.630 477.961
459.712 441.882 424.470 407.474 390.893 374.724 358.966 343.618
328.675 314.137 300.000 286.262 272.919 259.969 247.409 235.234
223.441 212.028 200.989 190.320 180.018 170.078 160.496 151.266
142.385 133.846 125.646 117.777 110.237 103.017 96.114 89.520
83.231 77.240 71.540 66.125 60.990 56.126 51.528 47.188
43.100 39.257 35.650 32.274 29.121 26.183 23.453 20.922
18.585 16.432 14.456 12.649 11.004 9.512 8.165 6.957
5.878 4.920 4.077 3.340 2.701 2.153 1.687 1.297
9.750E-01 7.140E-01 5.060E-01 3.450E-01 2.240E-01 1.370E-01 7.700E-02 3.800E-02
1.600E-02 5.000E-03 5.000E-05
-1.
%%%
```

## C EXAMPLE INPUT FILE FOR GIFTSFRTE

Below is an example input file for GIFTSFRTE. In this example MM5 temperature, moisture and ozone profiles are in file 06122002\_2200\_2\_3.bin, condensate effective diameters and mixing ratios are in file 06122002\_2200\_2\_3.diams.bin. The regression coefficients for clear sky are in files located on the path

```
/home/jimd/PROJECT/GIFTS/DOC/VERIFY/experiment_v6.01/FAST/
```

as are, in this case, the cloud optical properties file `clouds.dat`. Output files are to be prefixed by `liq_`. Records 594 to 599 are to be processed for both channels (1 is LW, 2 is SMW). Effective diameters for liquid and ice phase particles are given flag values (2000,2000 which are flag values if above 1000 ( $\mu\text{m}$ )) which means that the effective particle sizes will be computed from the condensate profiles as indicated in Sections 4.3.1 and 4.3.2. The final line, (0,5) sets the verbosity to zero (minimal output to STDOUT) and instructs GIFTSFRTE to generate only wavenumber files and brightness temperature files — the final number is a bitmask to indicate output required; bit 0 for wavenumbers, bit 1 for radiances, bit 2 for brightness temperatures and bit 3 for column transmittances.

```
06122002_2200_2_3.bin
06122002_2200_2_3.diams.bin
/home/jimd/PROJECT/GIFTS/DOC/VERIFY/experiment_v6.01/FAST/
/home/jimd/PROJECT/GIFTS/DOC/VERIFY/experiment_v6.01/FAST/
liq_
594,599
1,2
2000,2000
0,5
```

## D EXAMPLE INPUT FILE FOR LBLDIS

Below is an example input file for LBLDIS. In this example, radiances generated for a nadir view over the spectral range 550.0 to 2400.0 wavenumbers at 0.01 stepsize. Radiances for six optical depths are computed, for a single layer cloud whose top is at 10.1249 km. The effective particle *radius*<sup>4</sup> is 50  $\mu\text{m}$  at 1000  $\text{cm}^{-1}$  (i.e. 10  $\mu\text{m}$ ). The gaseous optical depth files from LBLRTM are located in directory /abyss/Users/jimd/water\_cloud\_experiment\_v6.01/iatm6\_1013\_to\_103p018 and the single scattering property database files are listed with fully qualified path names. The surface skin temperature is 288.2 K and the surface emissivity is 1 at 100  $\text{cm}^{-1}$  linearly interpolated to 1 at 3000  $\text{cm}^{-1}$  (i.e. the emissivity is constant with value 1).

Header line

```
000.0
550.0 2400.0 0.01
6
1
2 10.1249 50.0 1000. 0.1 0.5 1.0 2.0 3.0 5.0
/abyss/Users/jimd/water_cloud_experiment_v6.01/iatm6_1013_to_103p018
3
/home/jimd/LBLDIS/single_scatter_properties/ssp_db.mie_wat.gamma_sigma_0p100
/home/jimd/LBLDIS/single_scatter_properties/ssp_db.mie_ice.gamma_sigma_0p100
/home/jimd/LBLDIS/single_scatter_properties/ssp_db.hex_ice.gamma.0p100
288.2
2
100 1
3000 1
```

---

<sup>4</sup>Some care is required since LBLDIS specifies particle single scattering properties by their phase, crystal habit (if appropriate) and effective *radius* whereas GIFTSFRTE specifies on phase and effective *diameter*.