Aerosol effect on Infrared remote sensing of atmospheric structure

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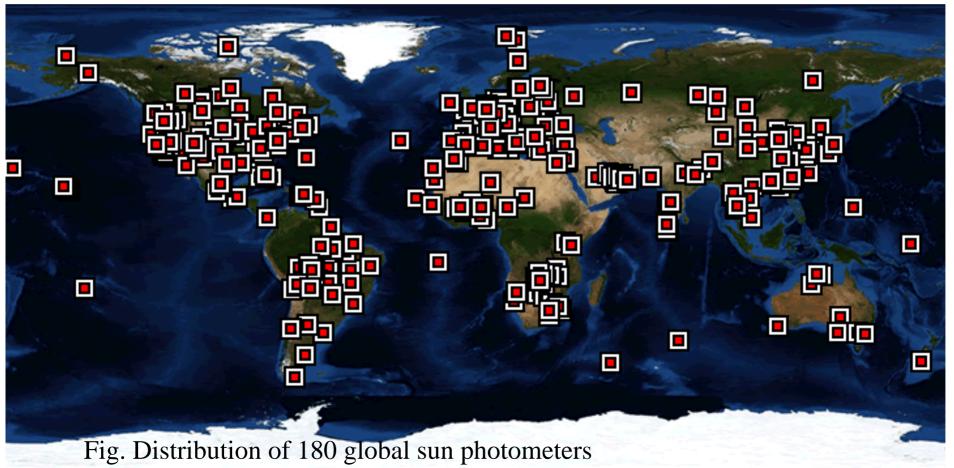
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Outline

1. Aerosol observations
2. Aerosol and optical effect
3. Effect and correction of aerosol on atmospheric retrieval

▶ 4. Conclusions

1.1 Ground-base Measurement

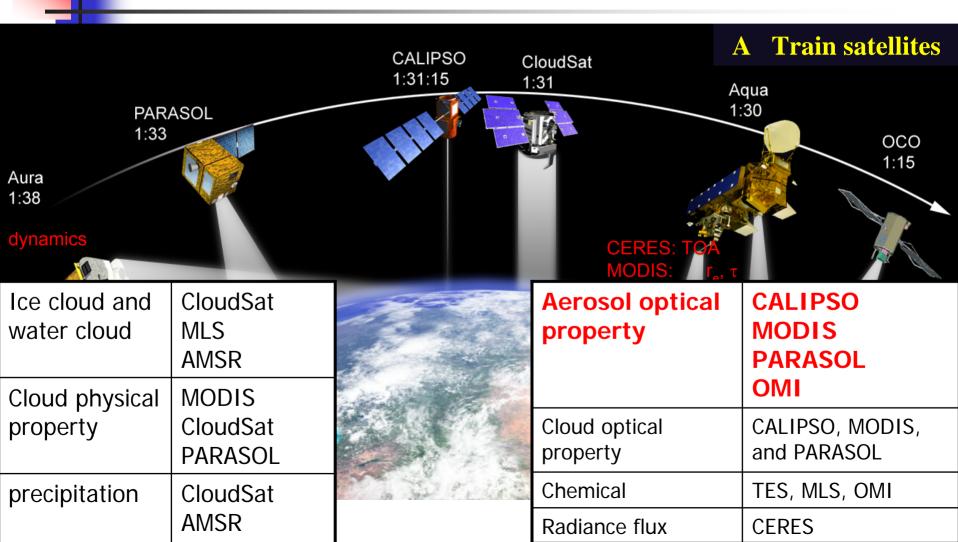


1.1 Ground-base Measurement



Fig. Automatic Sun Tracking Photometer CE-318 and laser radar CAMLTM CE 370-2

1.2 Airborne Measurement



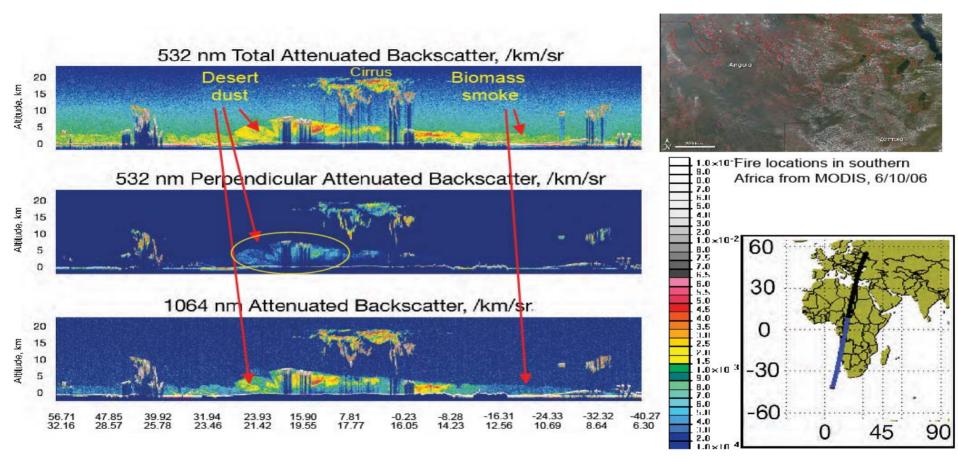
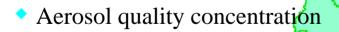


Fig. CALIPSO Level 1 profile data on 9 June 2006 provided by NASA Langley Dave Winker

1. Aerosol observation in China

1.3 sand storm forecast and warning service system



 Sun photometers, soil
 humidity, dry sedimentation, total particle Establish a national dust-storm forecast and warning service system by combining surface and satellite observations, surface physical-chemical parameters into the forecast model.

• tower Fig. Observed station distribution of Sun Photometers and other instrument

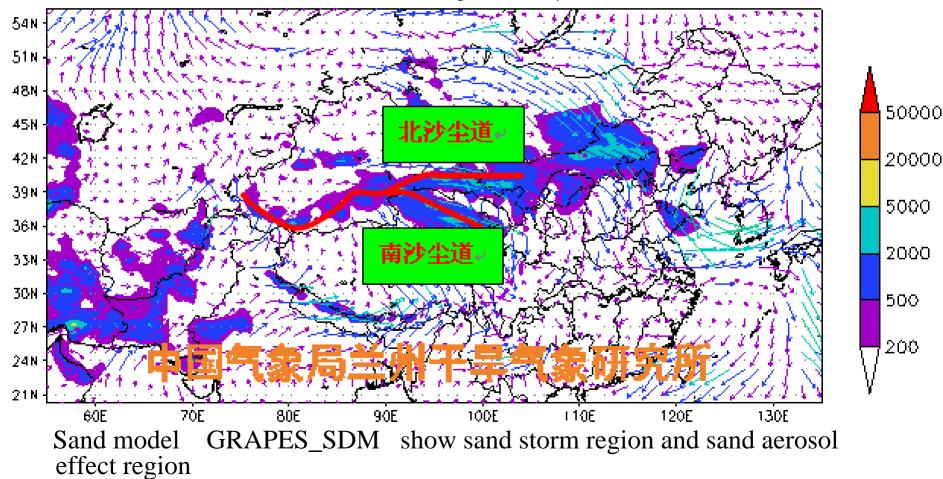
1. Aerosol observation in China

- 1.4 sand aerosol monitor from satellites
- EOS/MODIS 沙企监测图 2005年9月26日11日10日

中国气氛局兰州干旱气氛研究所 自由省气氛局卫星级成中心

1.4 sand storm forecast model

duststorm concentration(ug/m^3) at 2006030417



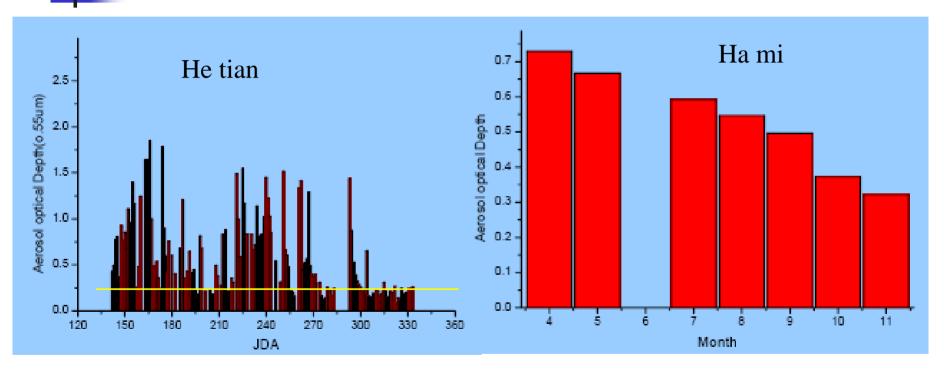
2.1 spatial distribution of aerosol optical depth (AOD)

Average AOD distribution in China show that The spatial distribution of Aerosol is inconsistent.

Fig. Average aerosol optical depth (AOD) In May 2002

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2.2 temporal variation of aerosol optical depth(AOD)



•AOD at He tian station is high in June , August and September, and low in July and winter, the low AOD is about 0.25.

•AOD at Ha mi station is high in April and decreased with month increased.

2.3 temporal variation of sand strom AOD

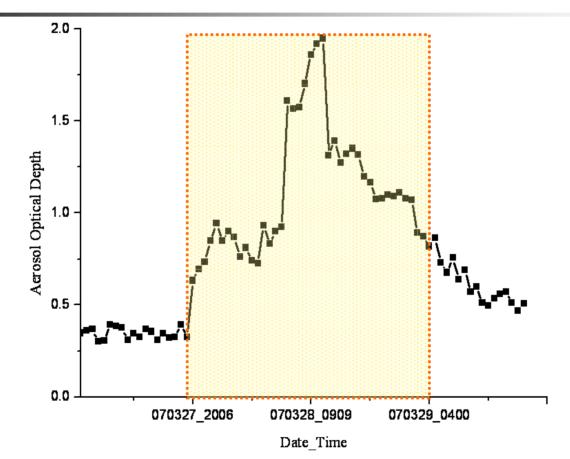


Fig. AOD variation when sand storm happened between 20:00 on 27 March 2007 and 4:00 on 29 March 2007

2.4 Aerosol change vertical structure of atmospheric temperature

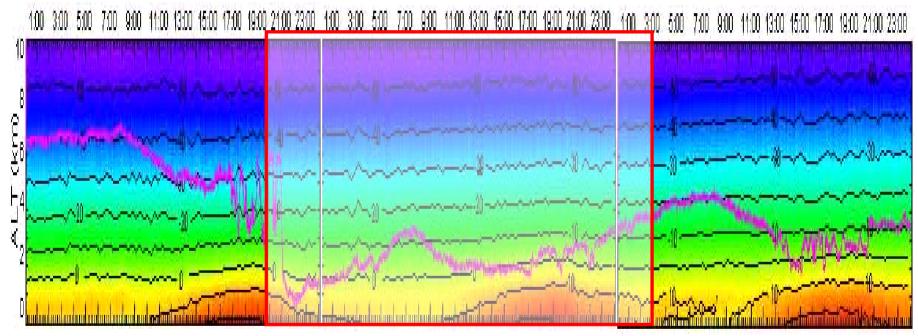


Fig. Temporal variable of atmospheric temperature from TP/WVP-3000 level

The reason is: though aerosol attenuate short-wave from Sun, it can hold back sensible heat transportation from surface, and increase downward long-wave radiance, the total effect is warming atmosphere .

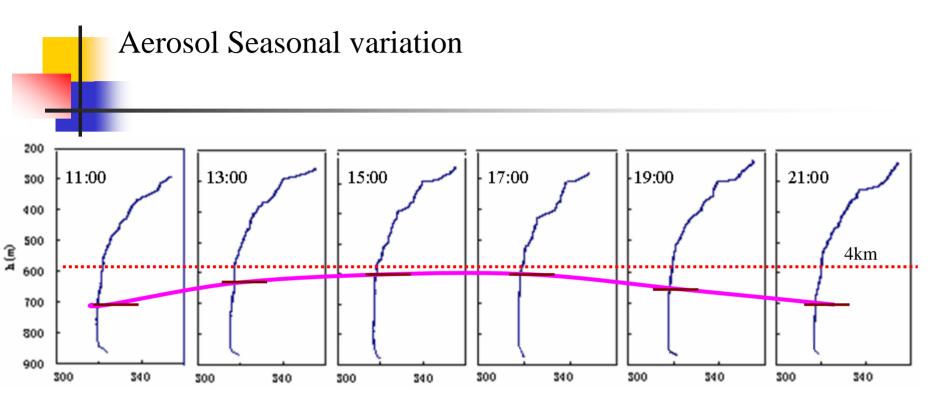
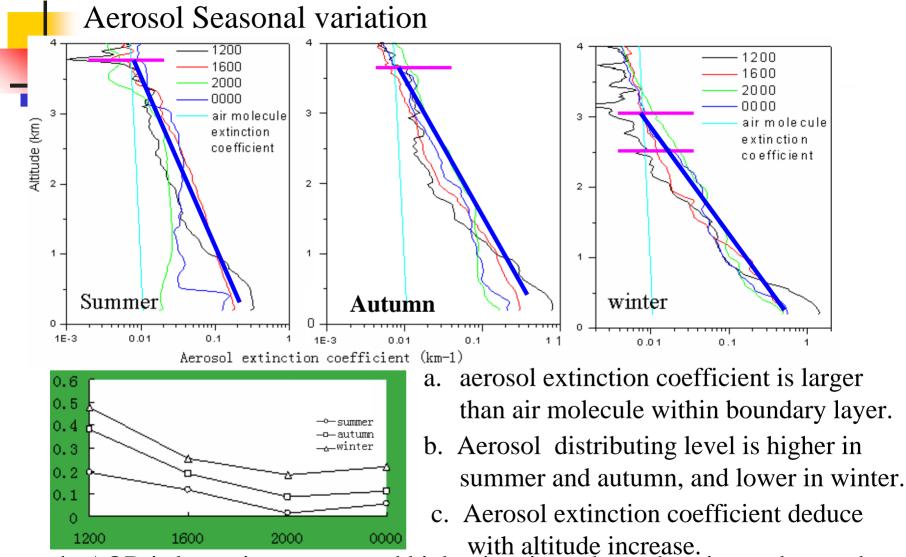
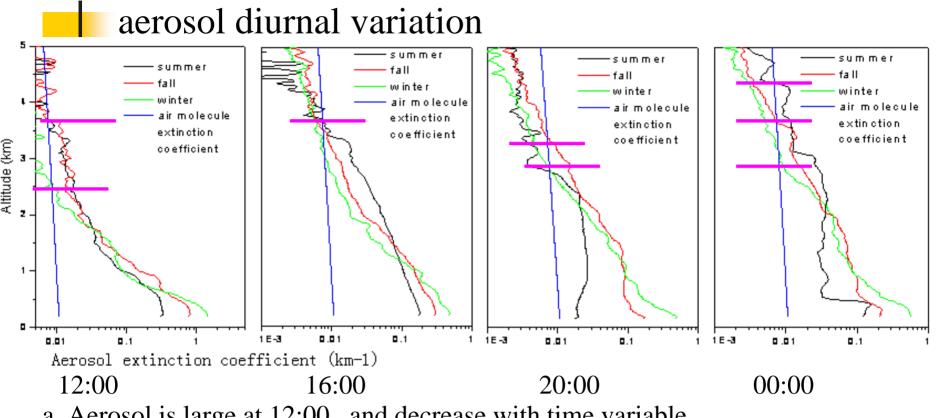


Fig. Potential temperature and the top of boundary layer diurnal variation on July 9 2006

The top of boundary layer is high at noon, and reach 600hpa or 4km level, which may effect aerosol vertical distribution



d. AOD is lower in summer and higher in winter due to burning carbon and biomass in winter.



a. Aerosol is large at 12:00, and decrease with time variable

b. There is a steady value at night, and the level is higher in summer and

lower in autumn and winter.

c. Aerosol level is higher between 12:00 and 16:00, lower at night, but it is increased late at night.

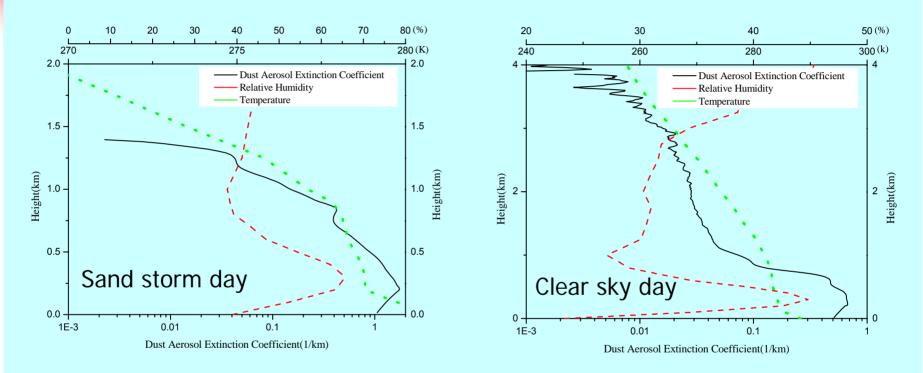


Fig. aerosol extinction coefficient change with atmospheric temperature and humidity from Profiling Radiometer TP/WVP-3000

Aerosol profile are in consistent with atmospheric temperature and relativity humidity profile



- Sand aerosol change the day-night variable of atmospheric temperature under boundary layer.
- Aerosol exist main in boundary layer, and the distributing is change with seasonal variable due to atmospheric environment change.
- Synchronous aerosol correction is necessary for remote sensing .

3.Effect and correction of aerosol on atmospheric temperature retrieval

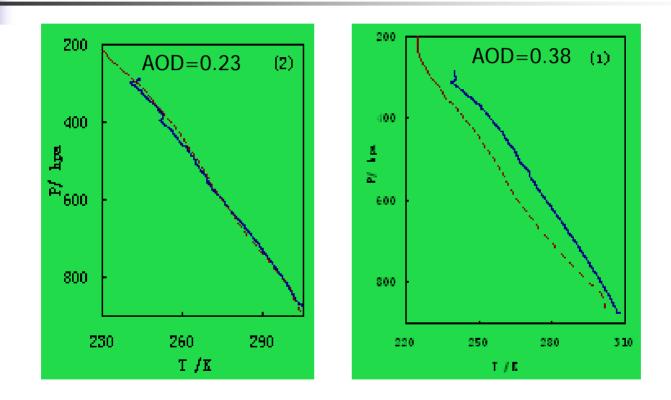
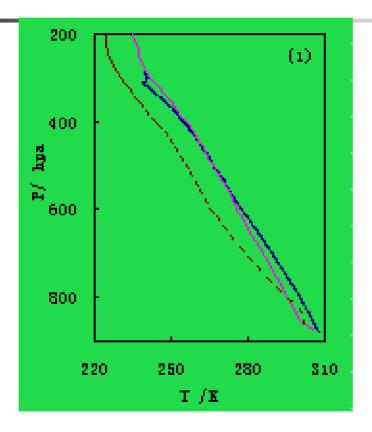


Fig. Statistical retrieval of atmospheric temperature from MODIS (dashed line indicates statistical retrieval, thick line is observed result)



the nonlinear iterative approach based on 1D-Var method used, the result is good for above the top of the boundary layer, but poor for within the boundary layer.

Fig. Physical retrieval of atmospheric temperature from MODIS (dashed line indicates statistical retrieval, thin line is physical retrieval value, thick line is observed result)

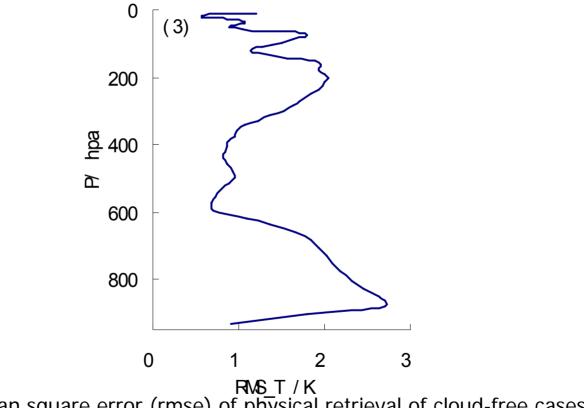


fig. the root mean square error (rmse) of physical retrieval of cloud-free cases

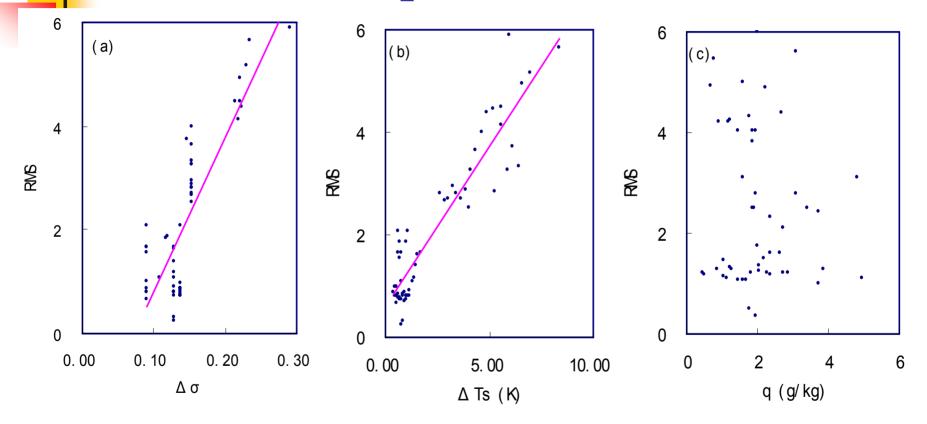
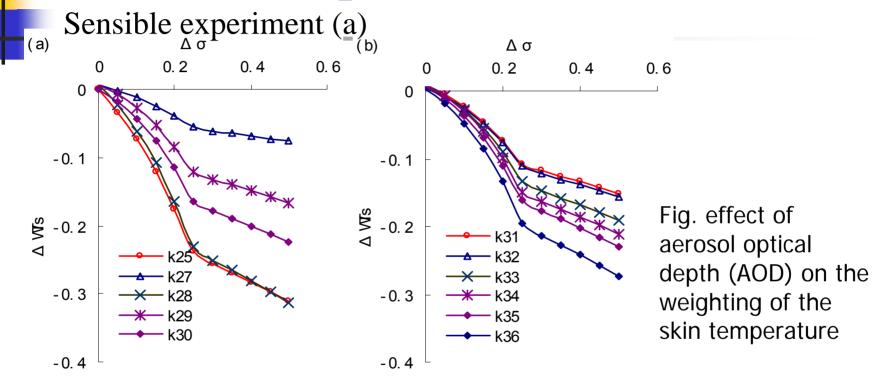


Fig. relations of temperature retrieval error at 2000m level with AOD increment (a), estimated error of skin temperature (b) and water vapor mixing ratio(c)



Weighting of IR band are decreased with AOD increase . If AOD is underestimated, atmospheric transmittance and weighting of skin and air temperature will be overestimated, and resulting in an underestimation of the retrieval temperature profile, ultimately leading to an increase in the absolute error.

Sensible experiment (b)

0.5 0.5 (b) (a) 25 0.4 27 0.4 28 33 29 0.3 34 0.3 30 SM ∇ 0.2 35 SM ∇ 0.2 36 0. 1 0.1 0 0 15 20 10 0 5 10 15 20 5 0 Δ Ts(K) Δ Ts(K)

Fig. effect of skin temperature estimated error on the weighting of the skin temperature

the weighting of the skin temperature is increased with increasing of skin temperature error, if skin temperature is overestimated, the weighting will be overestimated.

$$\tau_{a\lambda j} = \frac{p_j - p_t}{p_0 - p_t} \tau_{a\lambda}$$

$$\tau_{\lambda} = \tau_{ms\lambda} \cdot \tau_g \cdot \tau_o \cdot (\tau_r \cdot \tau_{a\lambda})$$

Considering the combined effects of continuum attenuation $\tau_{c\lambda}$ and molecular absorption $\tau_{ma\lambda}$ the spectral transmittance can be described as formula (2)

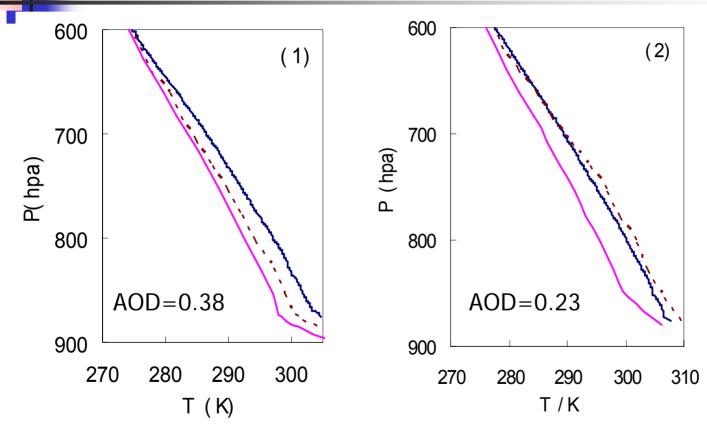


Fig. The contrast of no physical retrieval (thin line) with AOT correction profile (dash line) and observed value(thick line) of one case on 29 June 2006(left panel), another case with 0.23 AOD (middle panel)

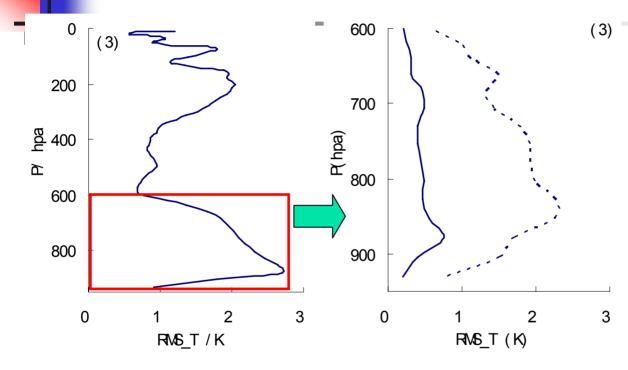


Fig. The right panel is the RMSE after AOD correction (dashed line) and the RMSE improvement (thick solid line) due to AOD correction from all 32 cases.

After AOD correction, the largest rmse of 32 cases is 2.34 K, the average rmse is 2.02K. Improvement rmse is in 0.15-0.58K, average improvement is 0.38K.

retrieval error : a) MODIS has low spectral resolution. b) better handling emissivity in retrieval should be studied; and c) the radiosonde error should also be considered.

4. Conclusions and future work

Aerosol leads to notable optical effect.

When aerosol effects at infrared spectral are considered, atmospheric transmittance can be better estimated. Moreover, the weighting of skin temperature and atmospheric temperature can be realistically estimated.

lessons learned from MODIS and the techniques tested on MODIS can be applied to both current (AIRS, IASI) and future (CrIS, HES) sounding instruments.

Considering aerosol size distribution in transfer model, to improve regional retrieval precision of IR sounding.



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